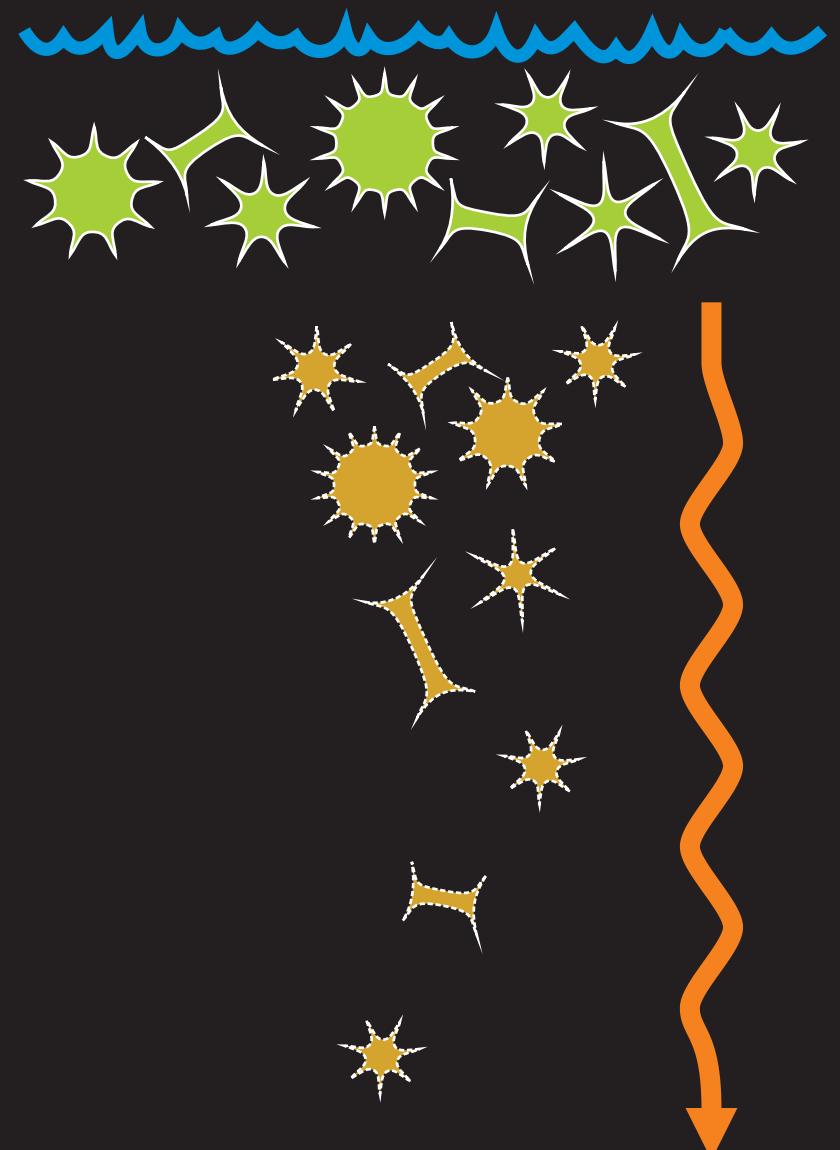
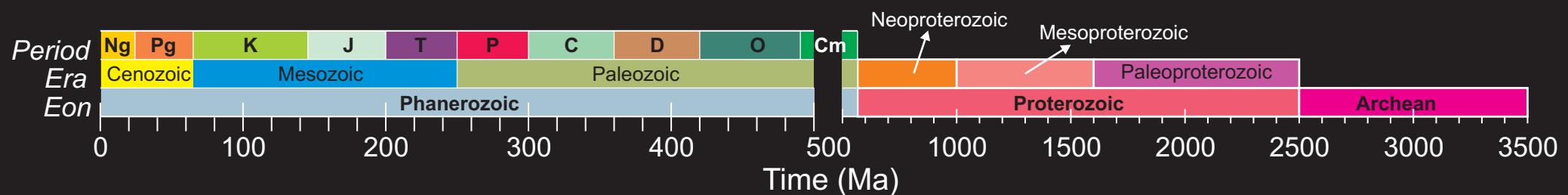


# Evolution of the Ocean's Biological Pump

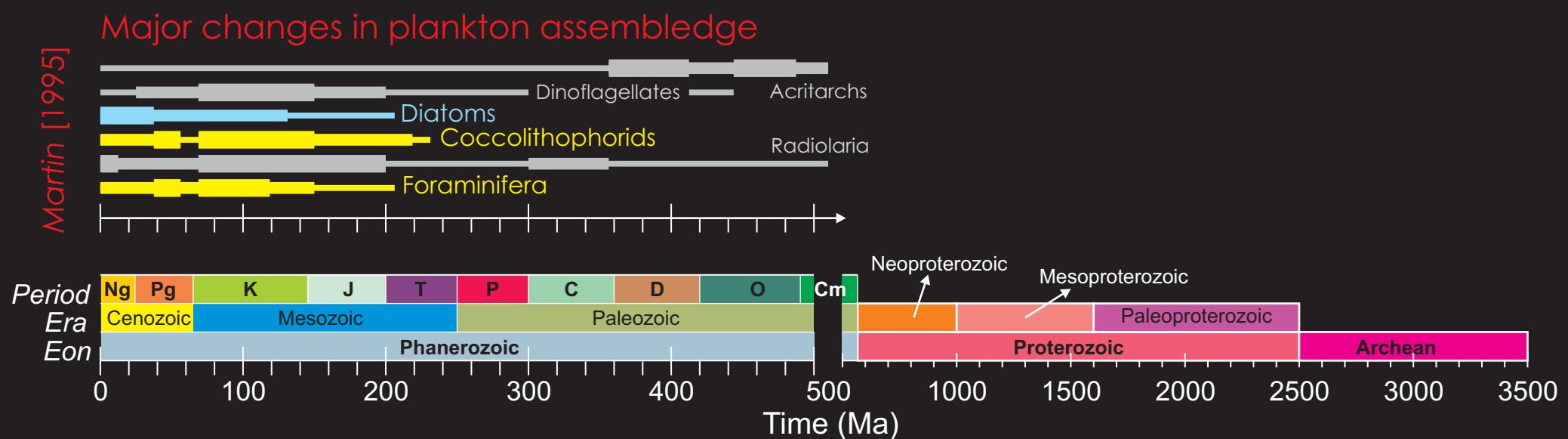
Andy Ridgwell



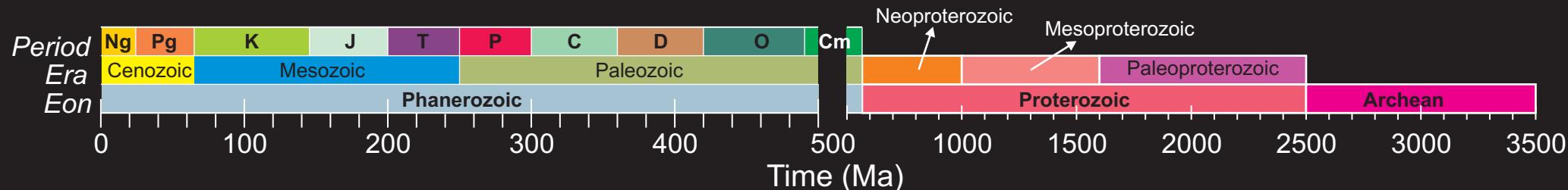
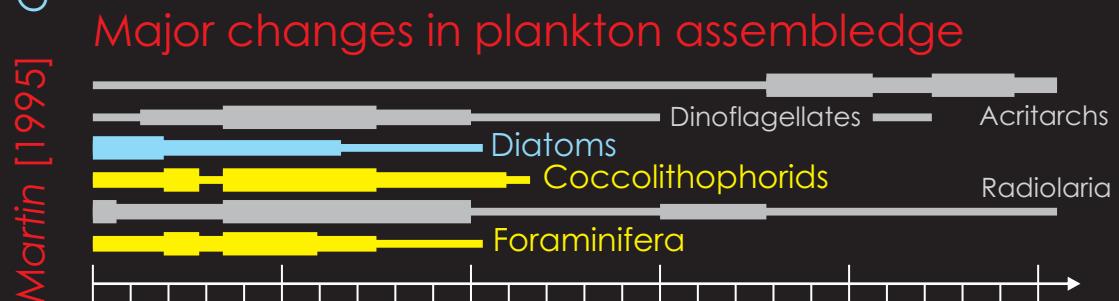
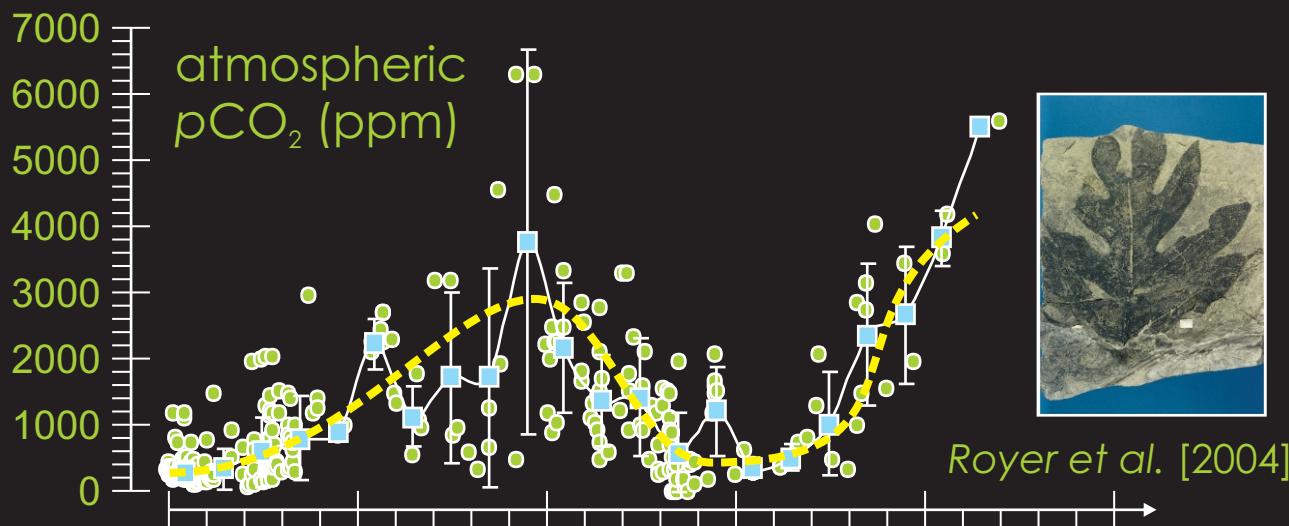
# *Evolution of the Biological Pump*



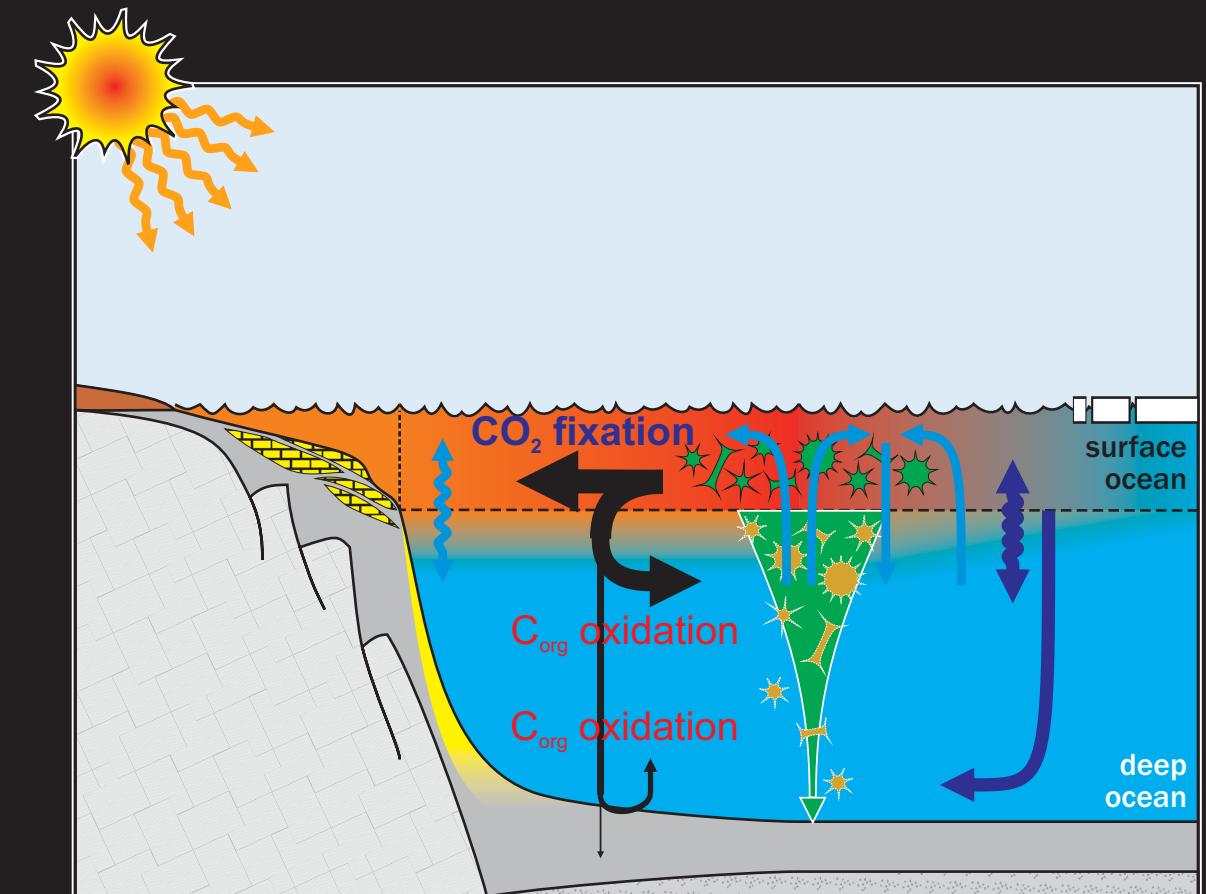
# *Evolution of the Biological Pump*



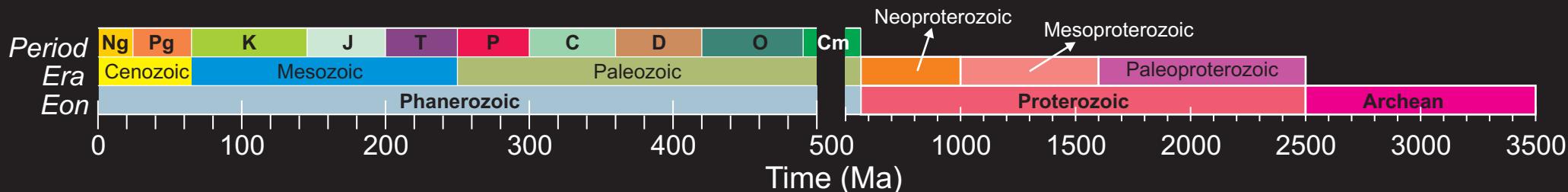
# Evolution of the Biological Pump



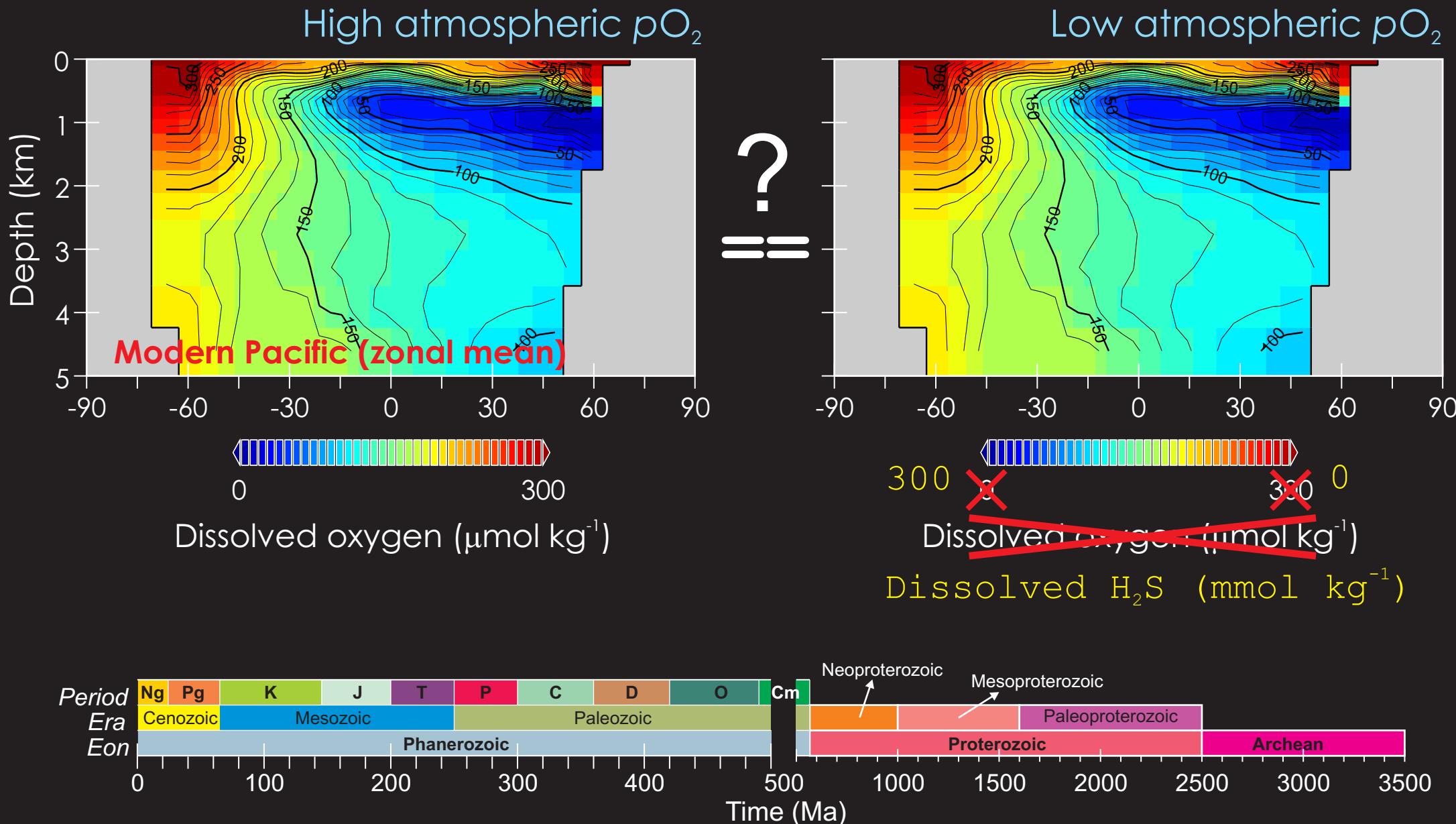
# Evolution of the Biological Pump



Both physical/geochemical and biological/ecological changes occurring through Earth history will affect the processes that govern the partitioning of carbon (and alkalinity) between the surface ocean (and hence atmosphere) and ocean interior, and conversely, oxygen.

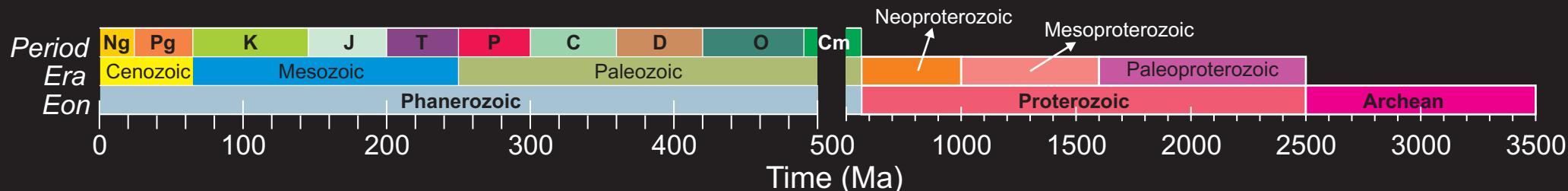


# Evolution of the Biological Pump



# *Evolution of the Biological Pump: TALK OUTLINE*

- (1) Interrogating the biological pump *in silico*.
  - (2) The fundamental importance ...  
or not ... of the advent of pelagic calcification  
and mineral ‘ballasting’ of particulate organic  
matter fluxes.
  - (3) Extinctions as a window onto the biological  
pump components.
  - (4) The fundamental importance ...  
or not ... of physical ocean changes and  
particularly warming.
  - (5) What came before the (vertical) particulate  
carbon pump?



# Evolution of the Biological Pump: in silico

```

! calculate carbonate alkalinity
loc_ALK_DIC = dum_ALK &
& - loc_H4BO4 - loc_OH - loc_HPO4 - 2.0*loc_PO4 - loc_H3SiO4 - loc_NH3 - loc_HS &
& + loc_H + loc_HSO4 + loc_HF + loc_H3PO4

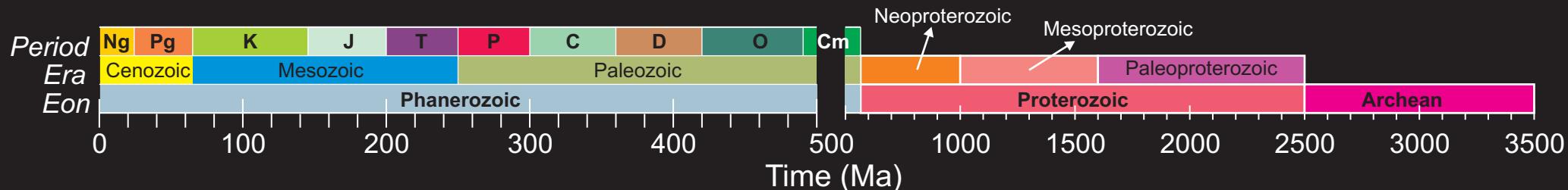
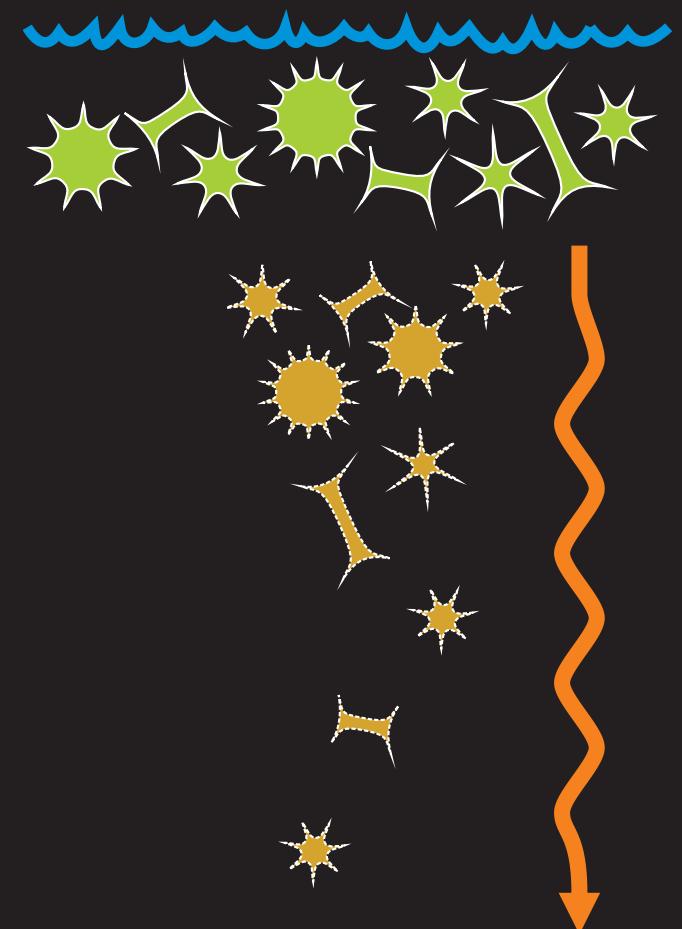
! estimate the partitioning between the aqueous carbonate species
loc_zed = ( &
& (4.0*loc_ALK_DIC + dum_DIC*dum_carbconst(icc_k) - 
loc_ALK_DIC*dum_carbconst(icc_k))**2 + &
& 4.0*(dum_carbconst(icc_k) - 4.0)*loc_ALK_DIC**2 &
& )**0.5 loc_conc_HCO3 = (dum_DIC*dum_carbconst(icc_k) - 
loc_zed) / (dum_carbconst(icc_k) - 4.0)

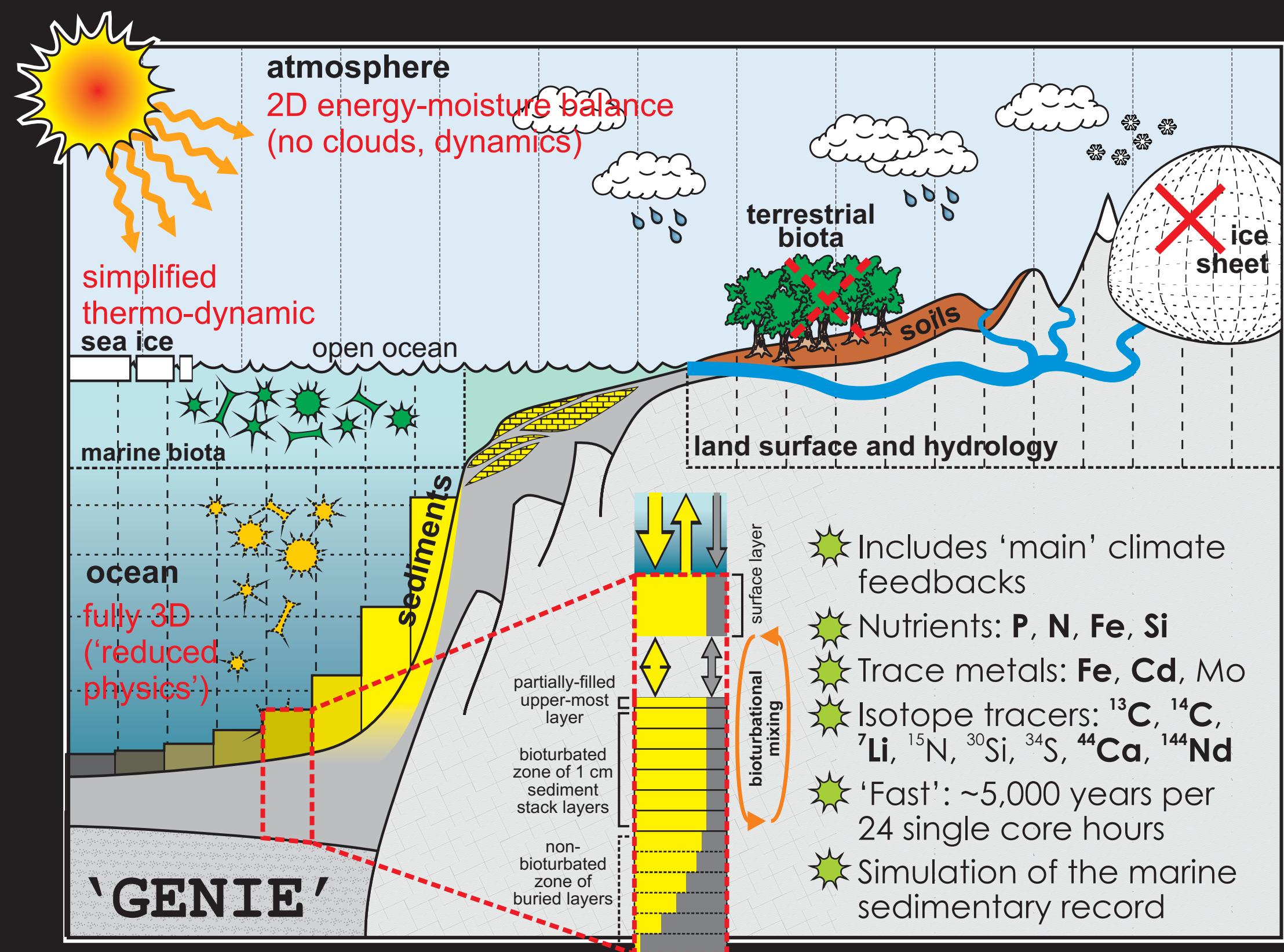
loc_conc_CO3 = &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) - dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& / (2.0*(dum_carbconst(icc_k) - 4.0))

loc_conc_CO2 = dum_DIC - loc_ALK_DIC + &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) - dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& / (2.0*(dum_carbconst(icc_k) - 4.0))

loc_H1 = dum_carbconst(icc_k1)*loc_conc_CO2/loc_conc_HCO3
loc_H2 = dum_carbconst(icc_k2)*loc_conc_HCO3/loc_conc_CO3

```





# cGENIE UCR 2014 version: README



Andy Ridgwell

March 5, 2014

1. To get an exact (read-only) copy of the ('mu~~n~~n' development branch)cGENIE source code used for the UCR presentation – in linux, (ideally from your home directory) type:

```
svn co https://svn.ggy.bris.ac.uk/subversion/genie/tags/cgenie.UCR2014  
--username=genie-user cgenie.muffin
```

NOTE: All this must be typed continuously on ONE LINE, with a S P A C E before '--username', and before 'cgenie'. You will be asked for a password – it is g3n1e-user.

2. You need to set a couple of environment variables – the compiler name, netCDF library name, and netCDF path. These are specified in the file user.mak (genie-main directory). If the cgenie code tree (cgenie.muffin ) and output directory (cgenie\_output) are installed anywhere other than in your account HOME directory, paths specifying this will have to be edited in: user.mak and user.sh (genie-main directory). Installing the model code under the default directory name (cgenie.mu~~n~~n) in your HOME directory is hence by far the simplest and avoids incurring additional/unnecessary pain (configuration complexity) ...

You will also need to have installed or linked to an appropriate FORTRAN compiler and netCDF library (built with the same FORTRAN compiler). The GNU FORT RAN compiler (gfort) version 4.4.4 or later is recommended. The netCDF version needs to be 4.0 (more recent versions require a little work-around, not documented here ...).

3. To test the code installation – change directory to cgenie.muffin/genie-main and type:

```
make testbiogem
```

This compiles a carbon cycle enabled configuration of GENIE and runs a short test, comparing the results against those of a pre-run experiment (also downloaded alongside the model source code). It serves to check that you have the software environment correctly configured. If you are unsuccessful here ... double-check the software and directory environment settings in user.mak (or user.sh) and for a netCDF error, check the value of the NETCDFDIR environment variable. (Refer to the User Manual for addition fault-finding tips.) If environment variables are changed: before re-trying the test, you will need to type:

```
make cleanall
```

That is for the basic installation. To run the model it is a simple matter of calling the 'runmuffin.sh ' shell script from genie-main and supplying a couple of parameter values, e.g.:

```
./runmuffin.sh cgenie.eb_go_gs_ac_bg.worjh2.ANT / EXAMPLE.worjh2.Caoetal2009.SPIN 10000
```

Refer to the cGENIE User manual for more information regarding installing, running, and analyzing model output, and cGENIE Examples for more information on this specific example.<sup>1</sup> Also read the cGENIE README

Highly recommended ... is in order to have a working appreciation of the structure of the model and output, plus the format of the model output and how to visualize it – to read through:

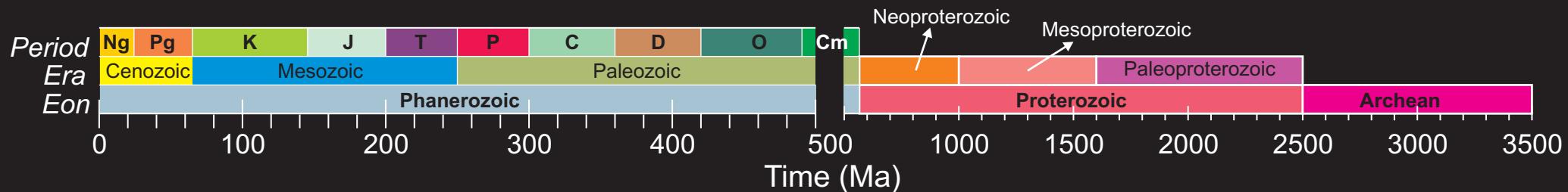
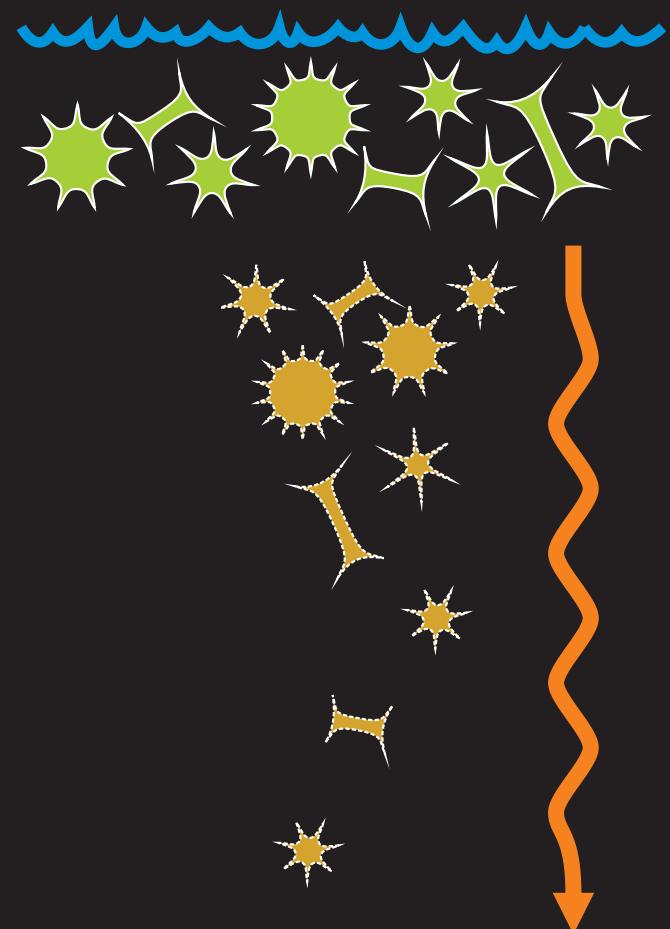
[http://www.seao2.info/cgenie/labs/EC4.2013/GEOGM1110andM1404.2013-14.cGENIE\\_LAB.0000.pdf](http://www.seao2.info/cgenie/labs/EC4.2013/GEOGM1110andM1404.2013-14.cGENIE_LAB.0000.pdf)

(which serves as a basic introduction to the model and how to see it).

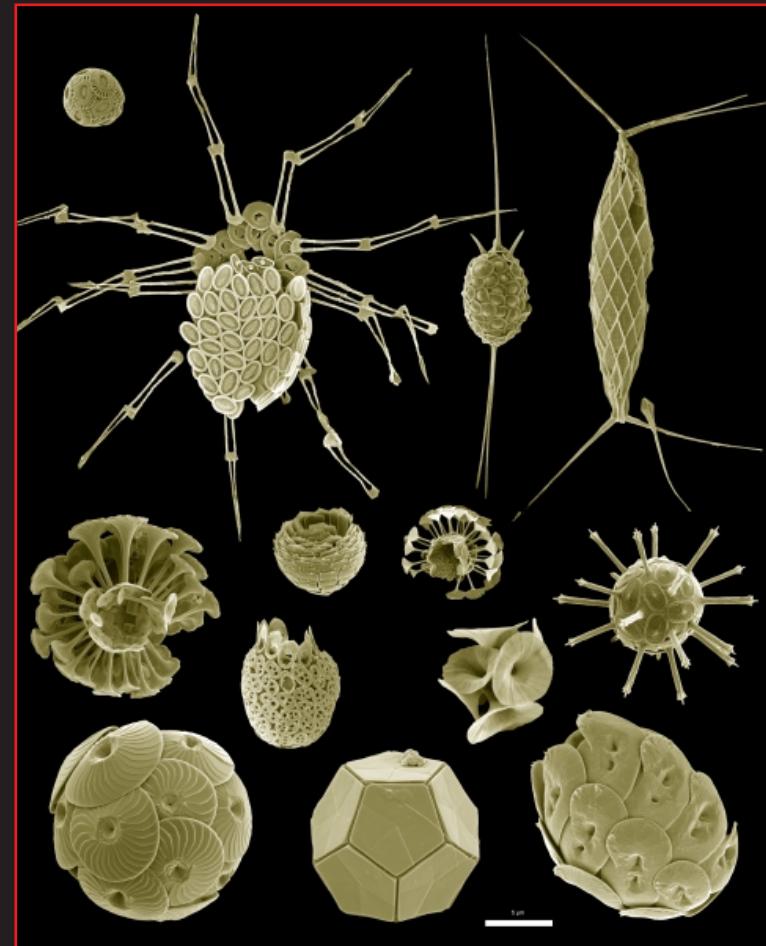
# *Evolution of the Biological Pump: in silico*



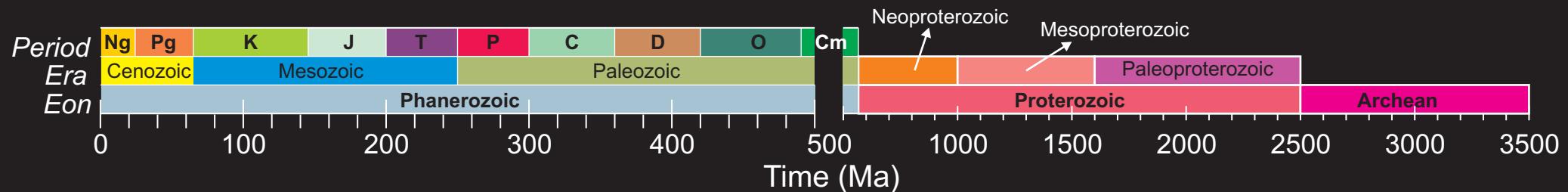
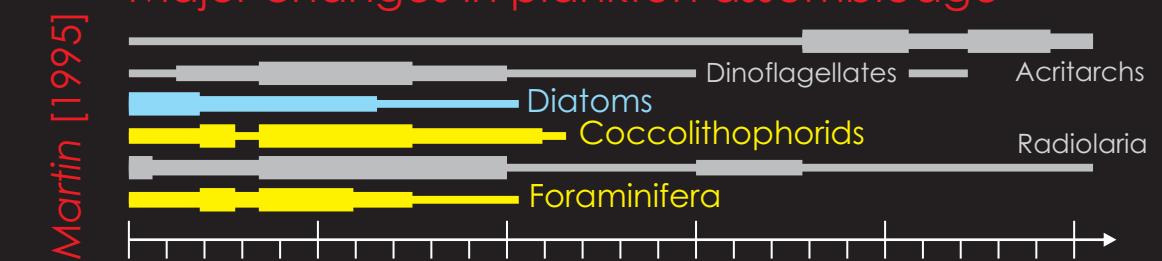
[tinyurl.com/kmjhe4s](http://tinyurl.com/kmjhe4s)



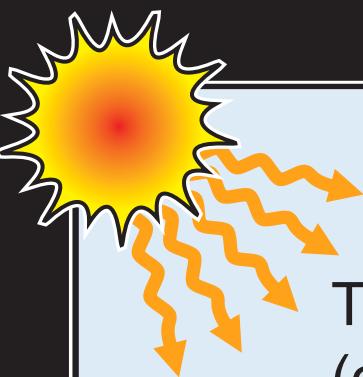
# *Evolution of the Biological Pump:* The Mesozoic planktic calcifier revolution



## Major changes in plankton assemblage



# Evolution of the Biological Pump



The stability of  $\text{CaCO}_3$   
(or thermodynamic favourability of precipitation)  
is defined by its 'saturation state':

$$\Omega = [\text{Ca}^{2+}] \times [\text{CO}_3^{2-}] / k$$

'Abiotic' precipitation rate  
(proportional to global burial)  
goes as:

$$r = f \times (\Omega - 1)^n$$

where  $n \sim 1.5\text{-}2.0$

low-temperature  
basaltic alteration



shallow water  
 $\text{CaCO}_3$  burial

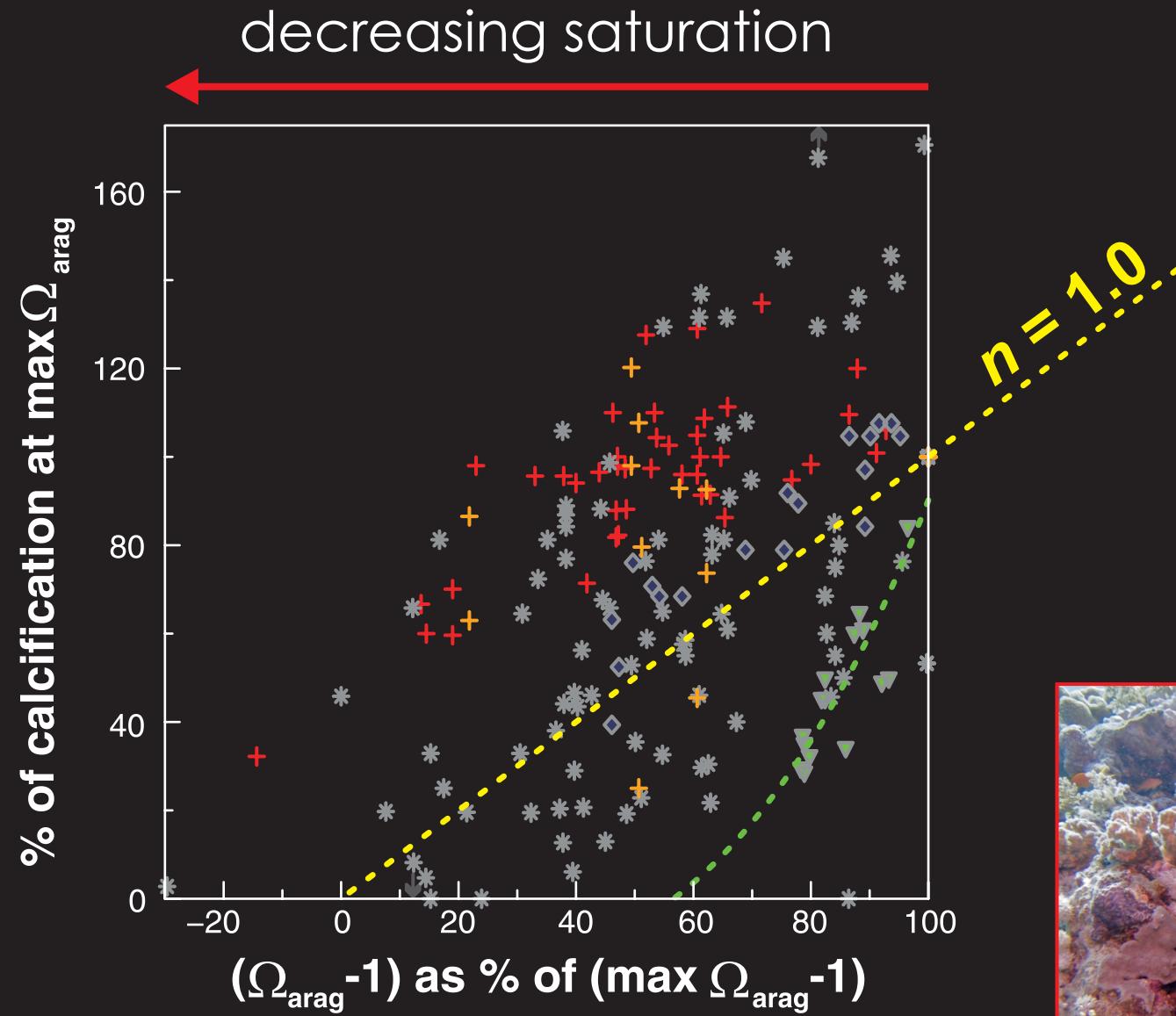


volcanism

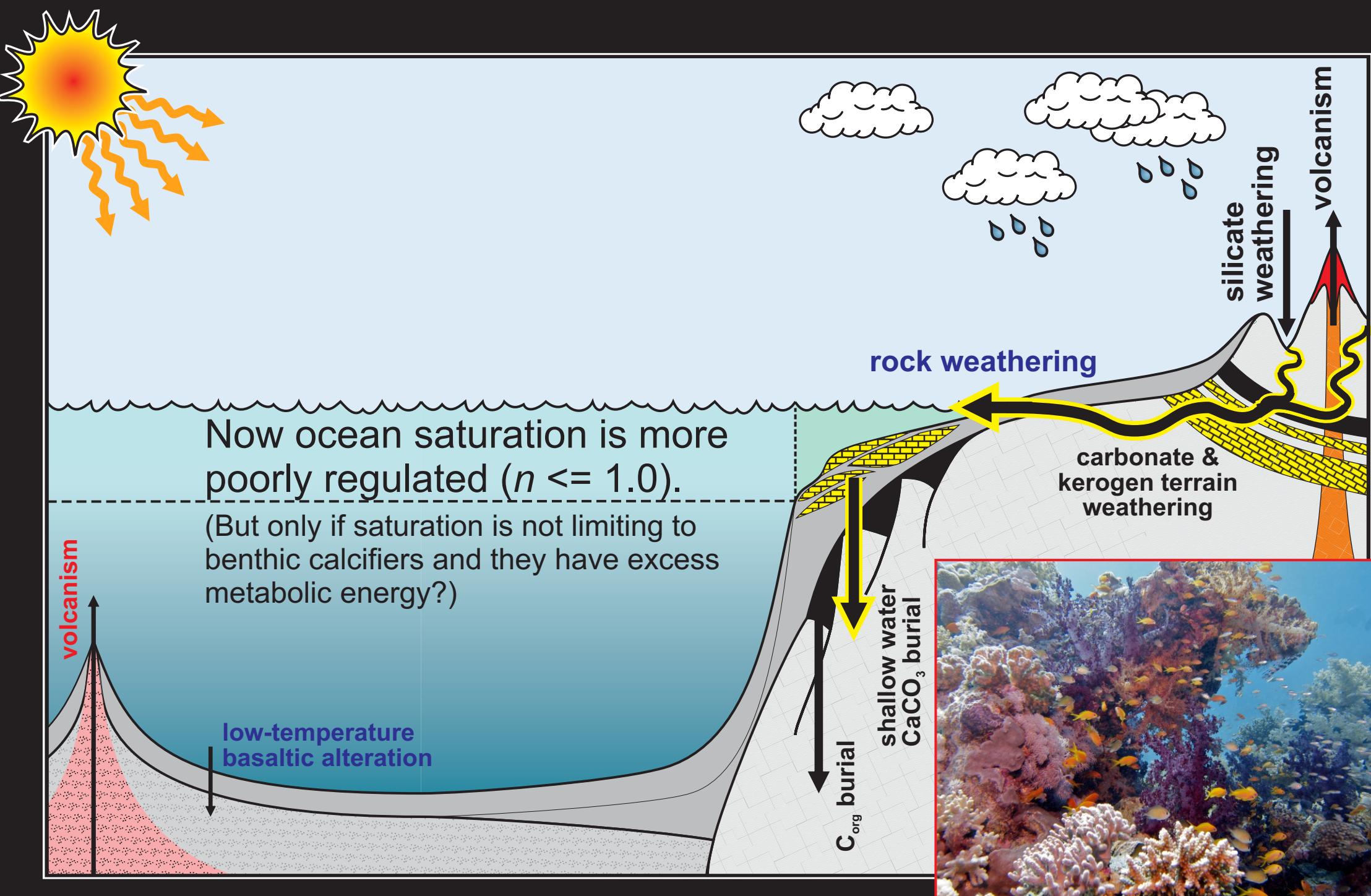
silicate  
weathering

# *Evolution of the Biological Pump*

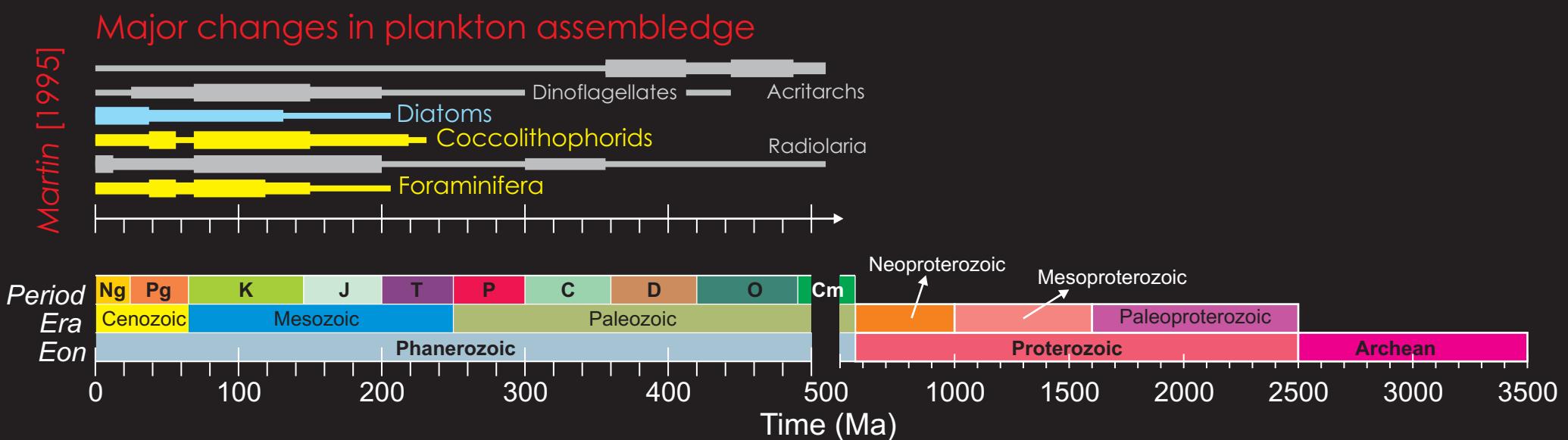
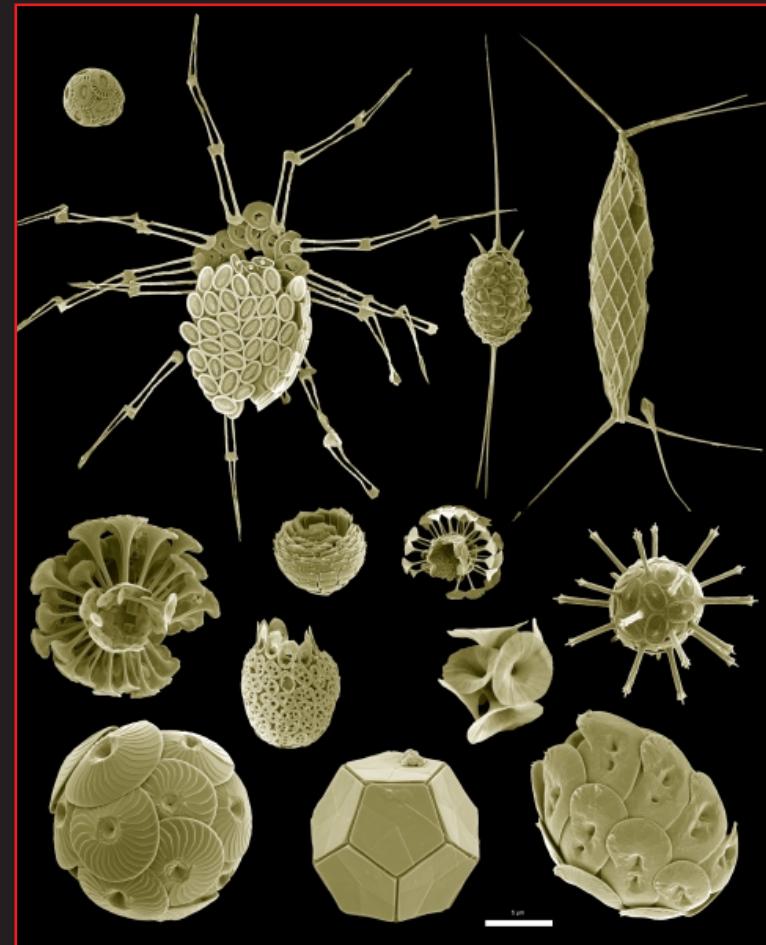
decreasing calcification rates  
(% compared to Preindustrial conditions)



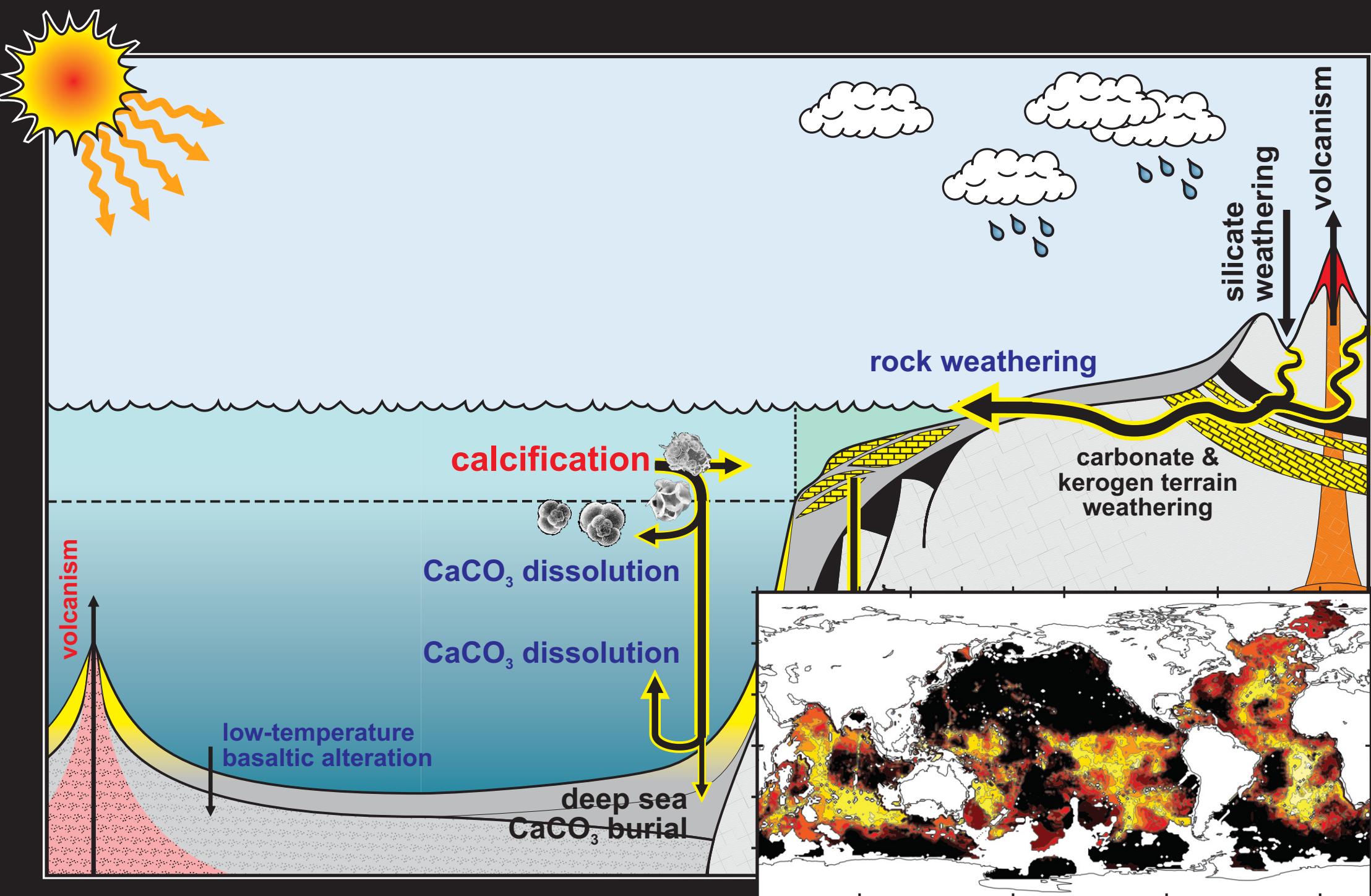
# Evolution of the Biological Pump



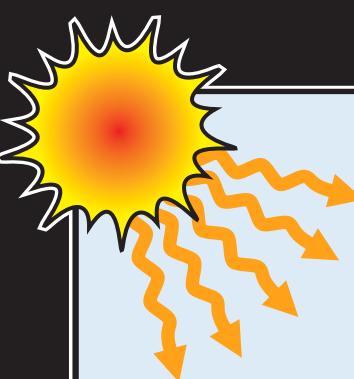
# *Evolution of the Biological Pump:* The Mesozoic planktic calcifier revolution



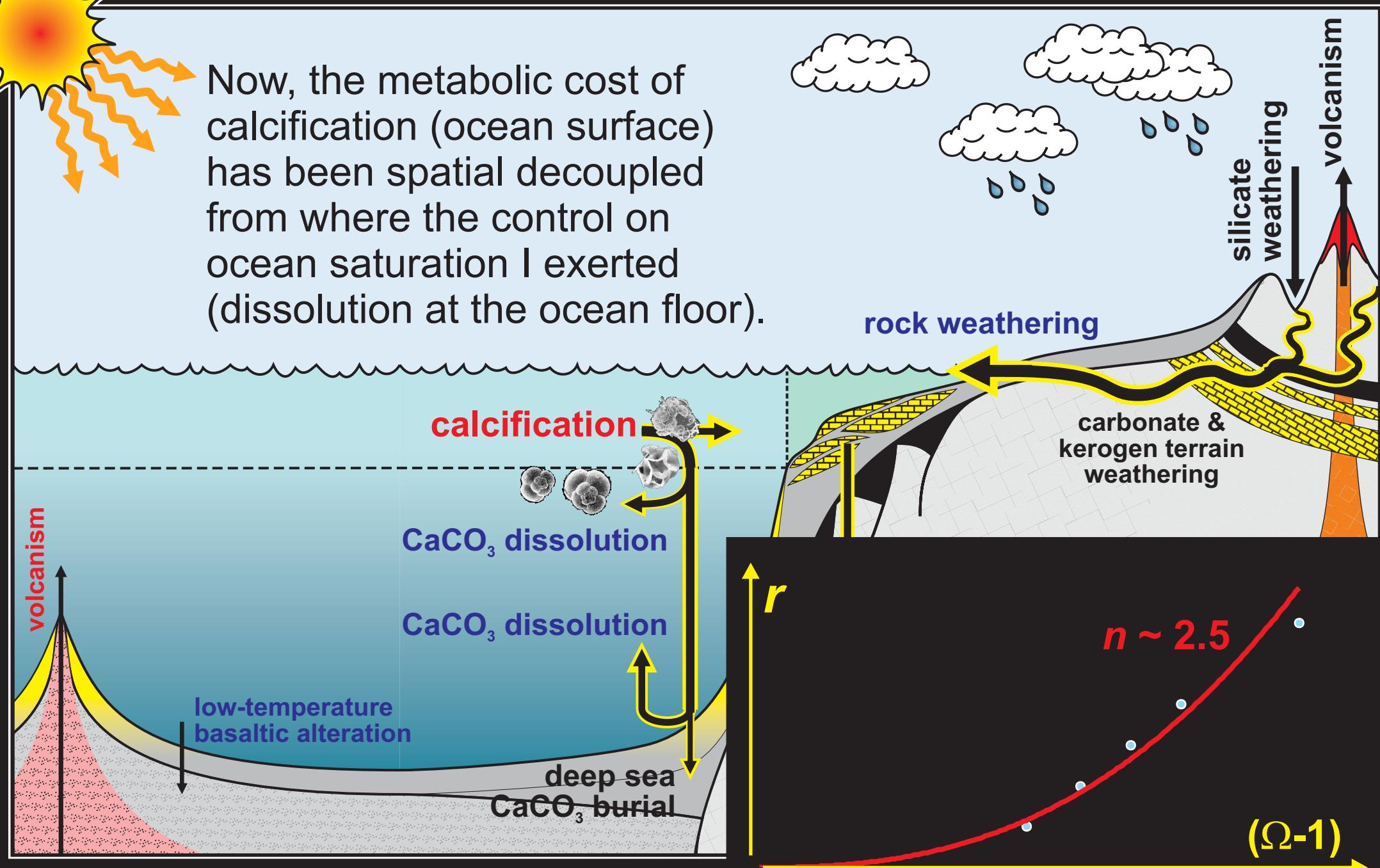
# Evolution of the Biological Pump

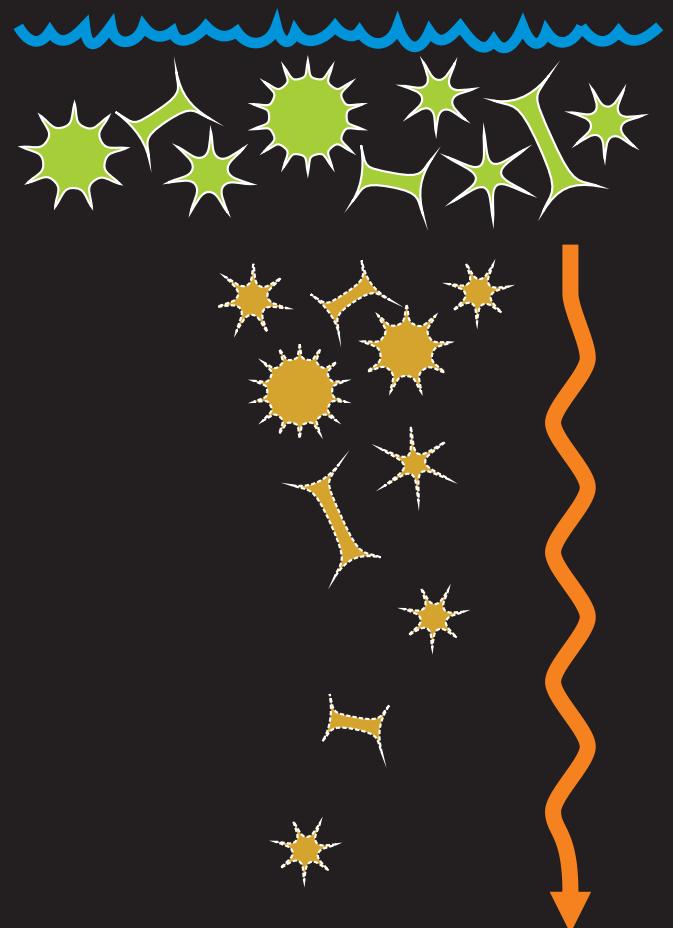


# Evolution of the Biological Pump

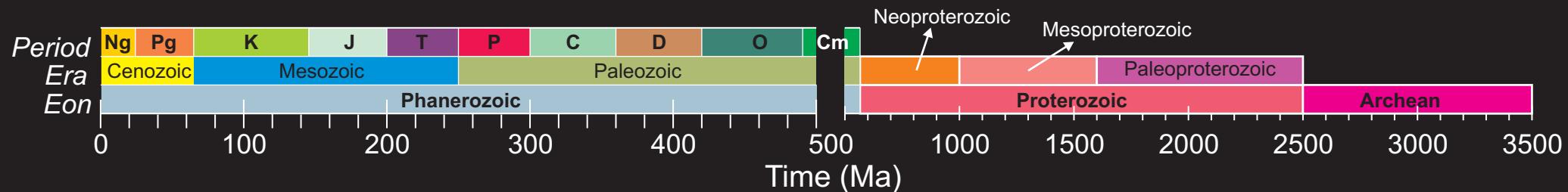
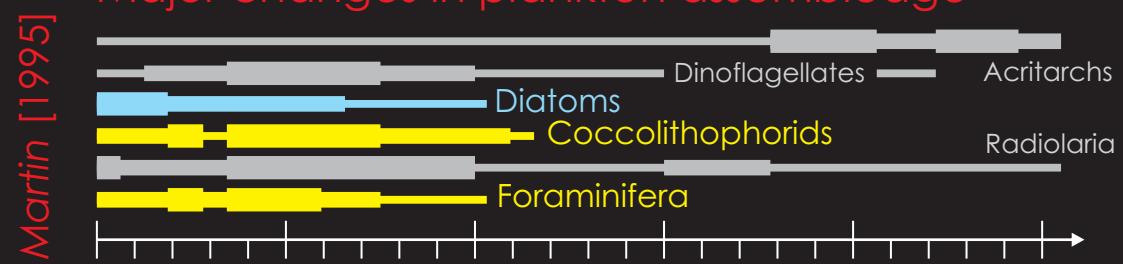


Now, the metabolic cost of calcification (ocean surface) has been spatially decoupled from where the control on ocean saturation I exerted (dissolution at the ocean floor).





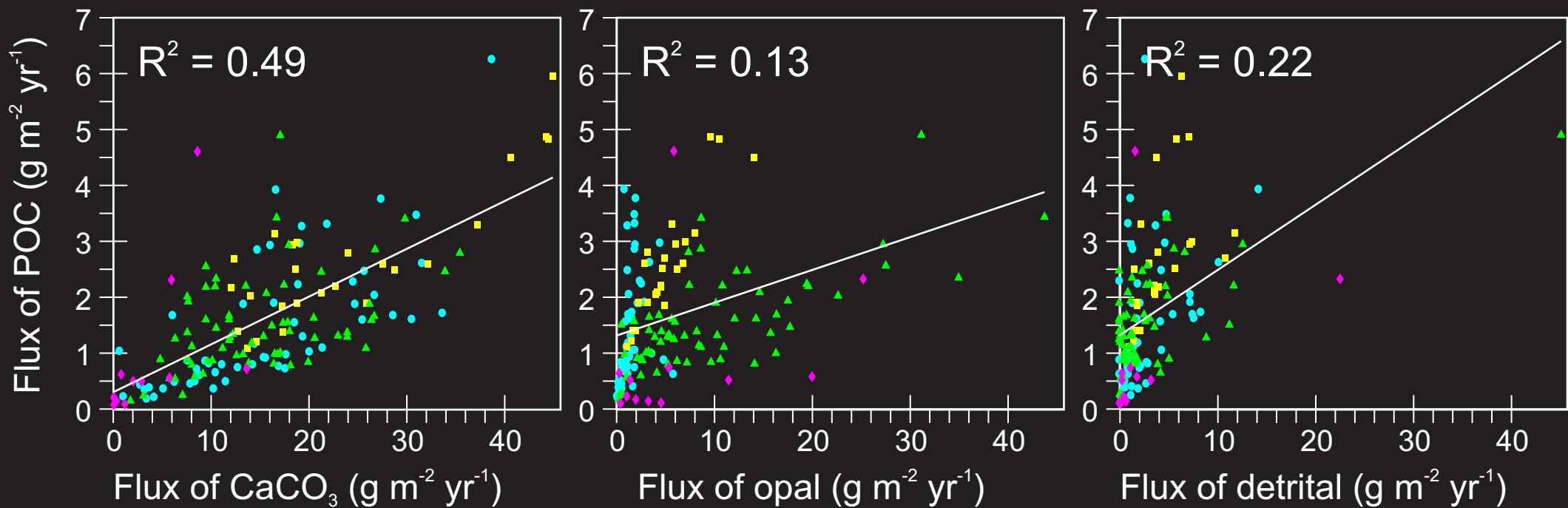
### Major changes in plankton assemblage



# *Evolution of the Biological Pump:* Planktic carbonate production and ‘ballasting’

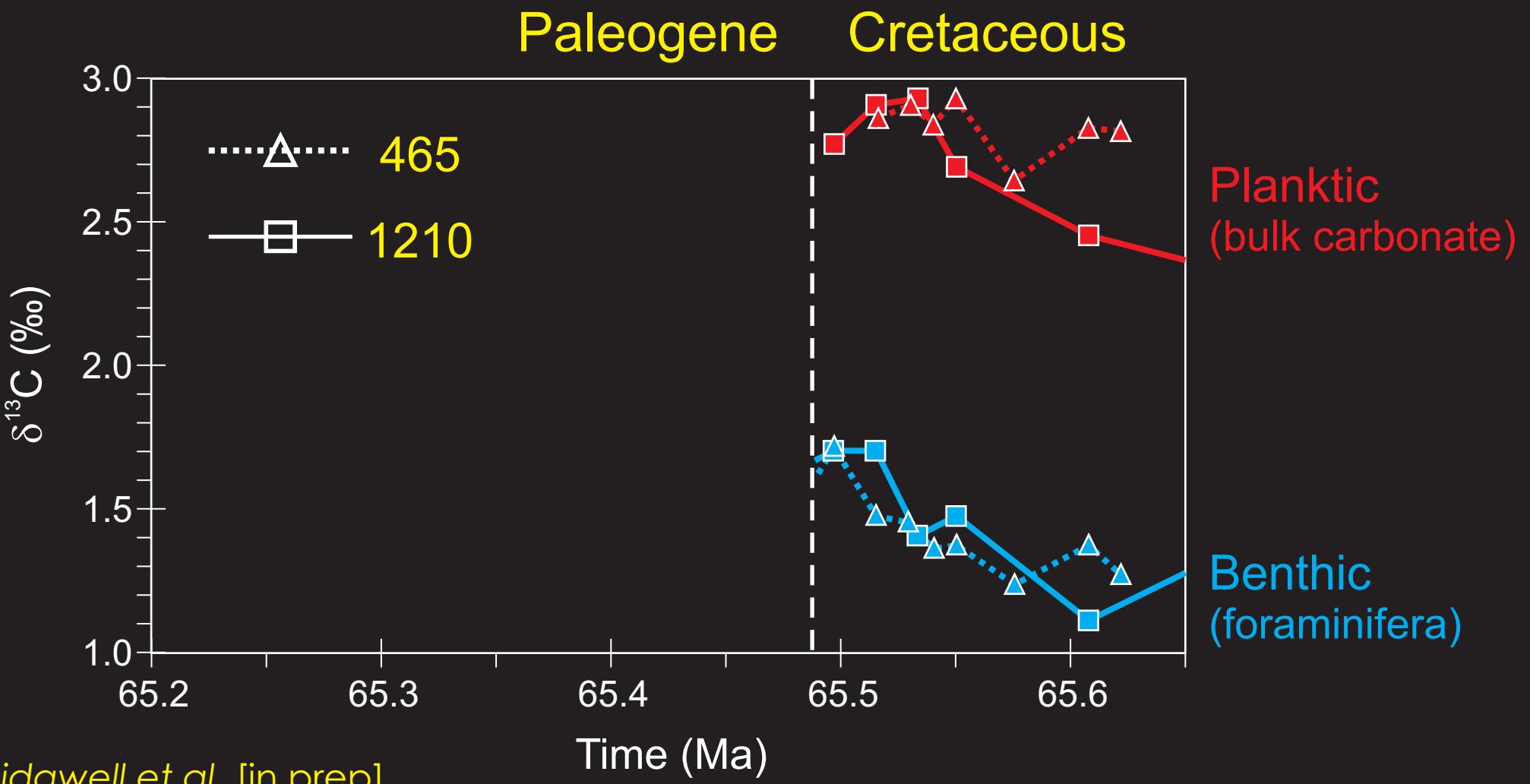
Compilation of sediment trap observations:  
depths  $\geq 2000$  m (to exclude hydrodynamically distorted  
fluxes and relationships) and differentiated by basin:  
cyan == Atl, yellow == Ind, green == Pac, magenta == SO.

[Wilson et al., 2012; GBC 26, doi:10.1029/2012GB004398]

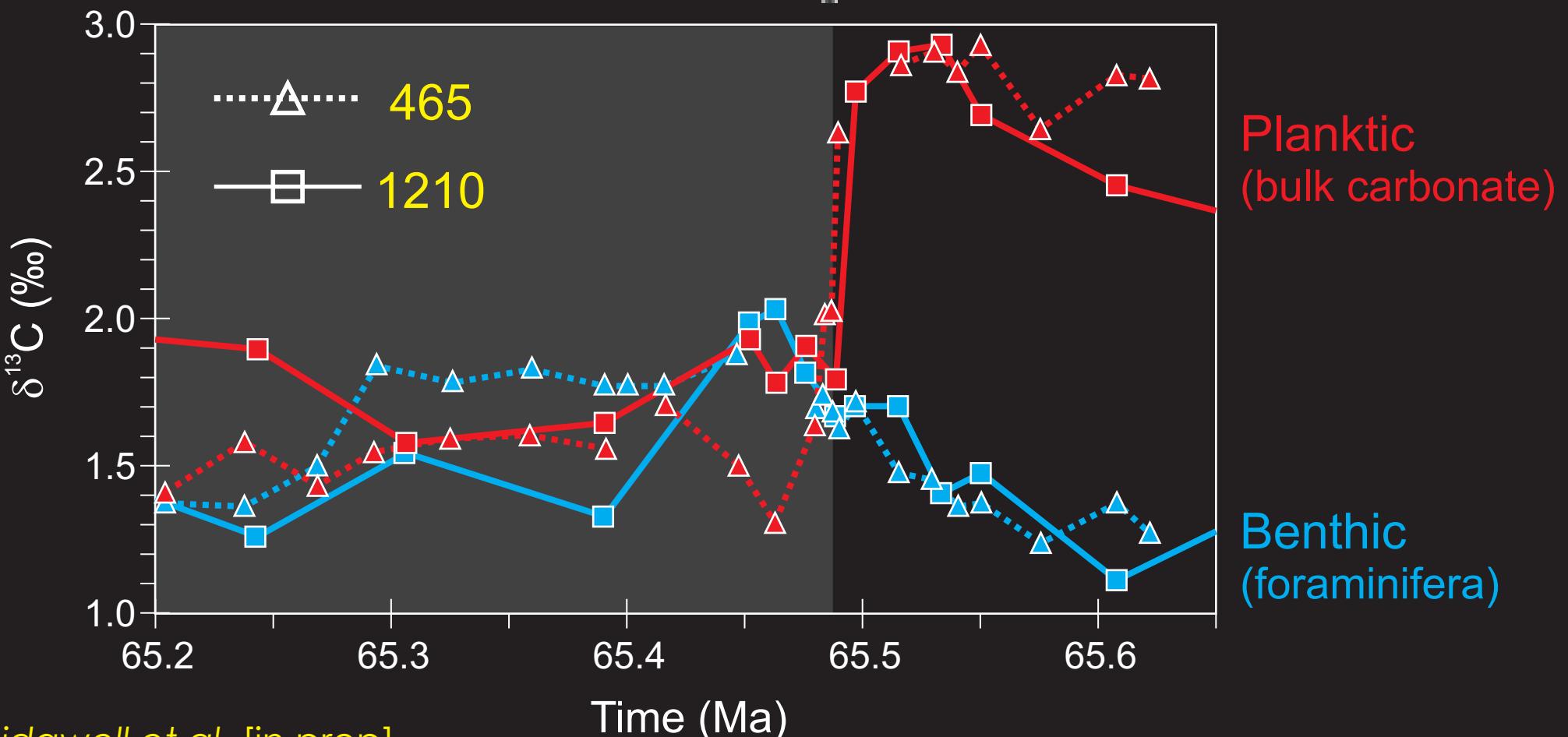


# *Evolution of the Biological Pump: ‘Hiccups’*

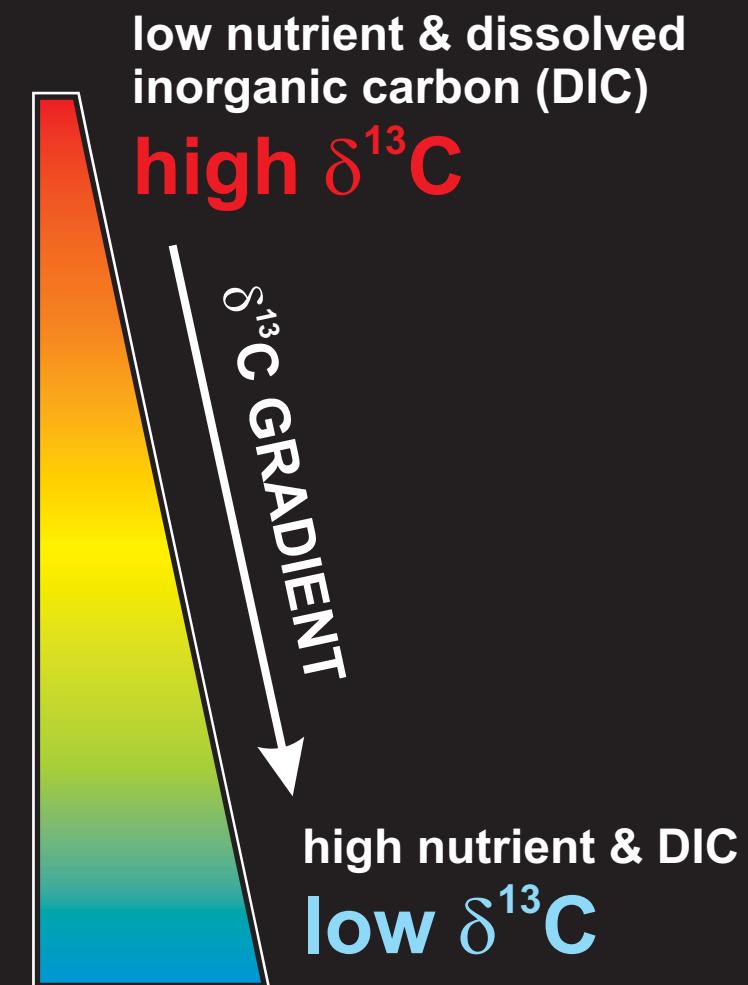
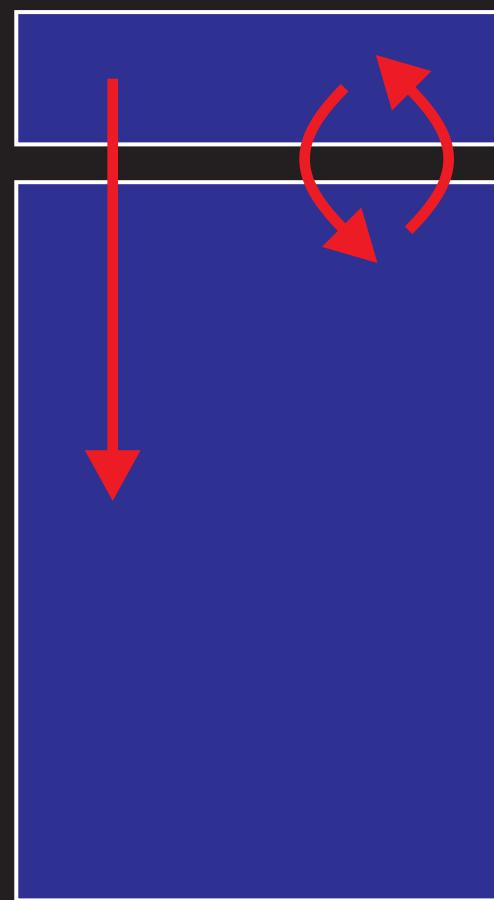
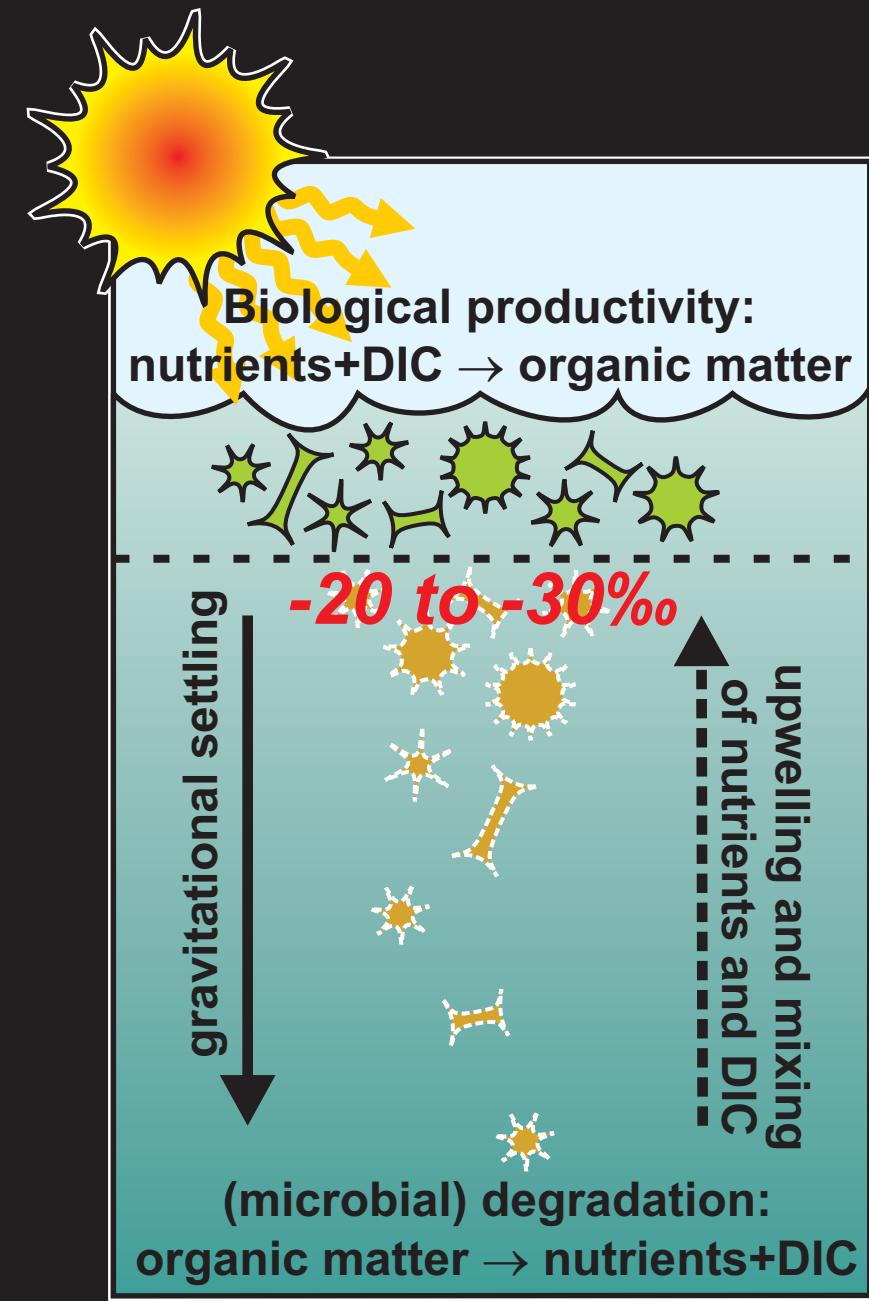
(temporary disruption or removal of one or more processes)



# *Evolution of the Biological Pump: ‘Hiccups’* (temporary disruption or removal of one or more processes)



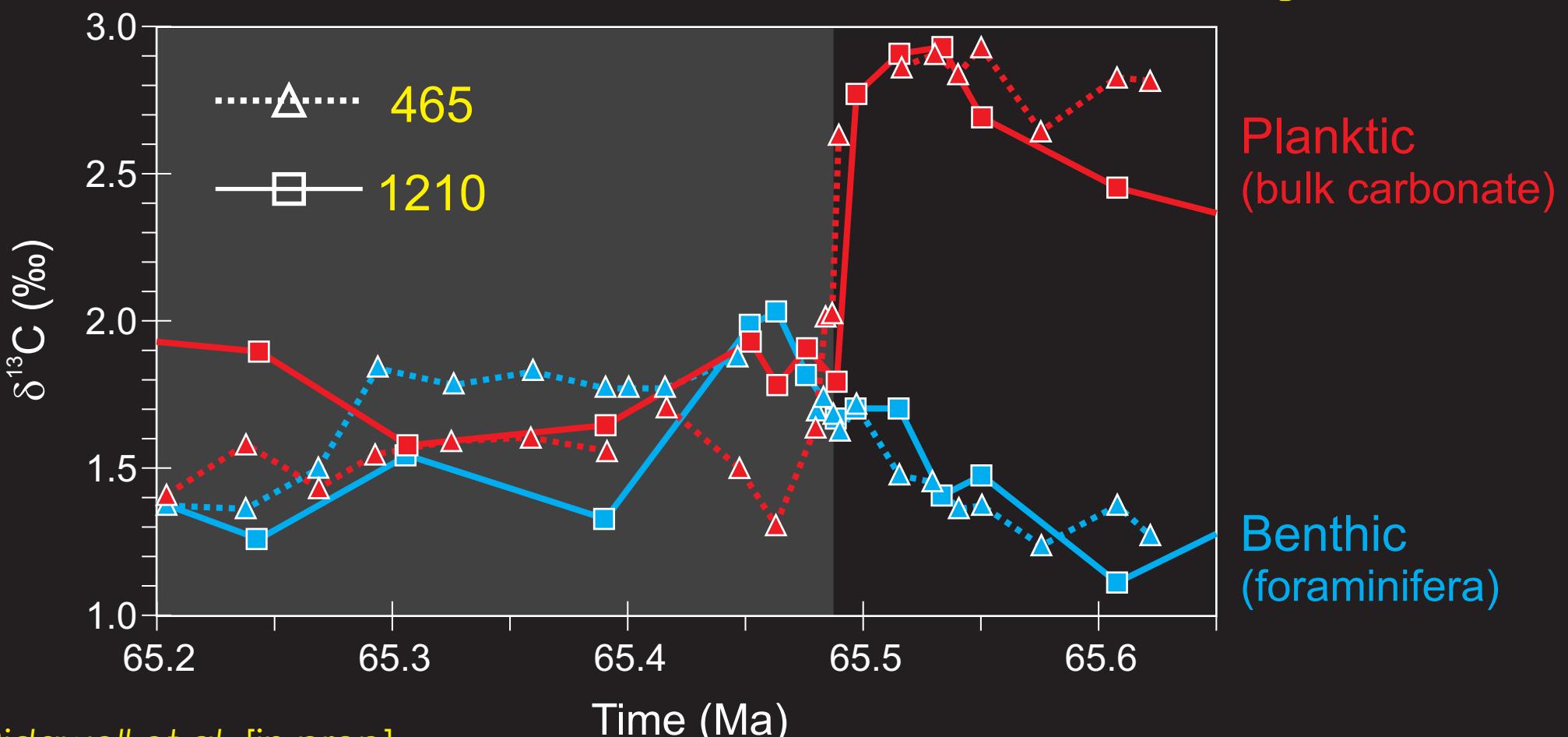
# Evolution of the Biological Pump: 'Hiccups'



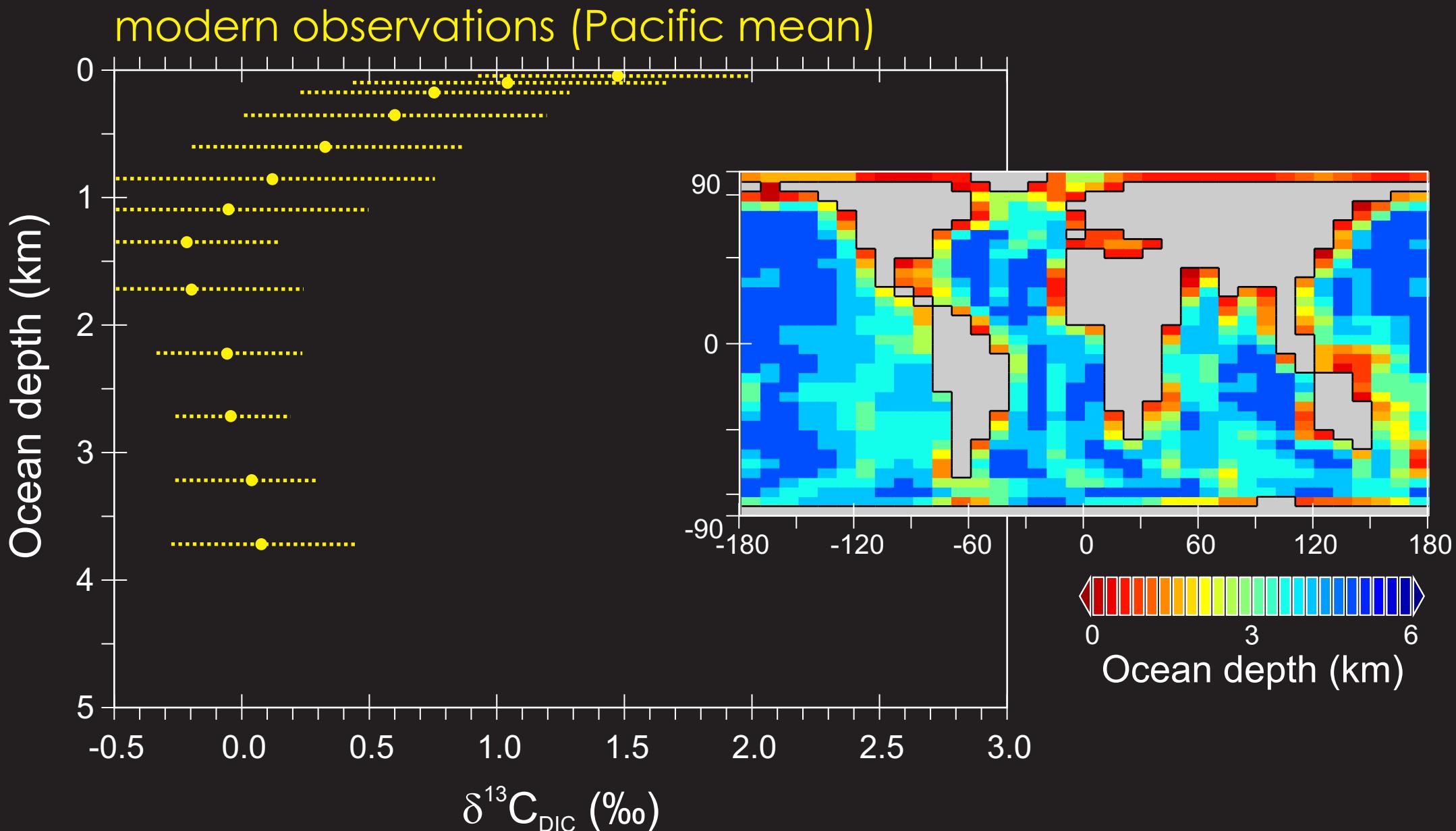
# *Evolution of the Biological Pump: ‘Hiccups’*

Severe extinction amongst  
calcifying plankton  
(and less interesting creatures  
such as dinosaurs etc.)

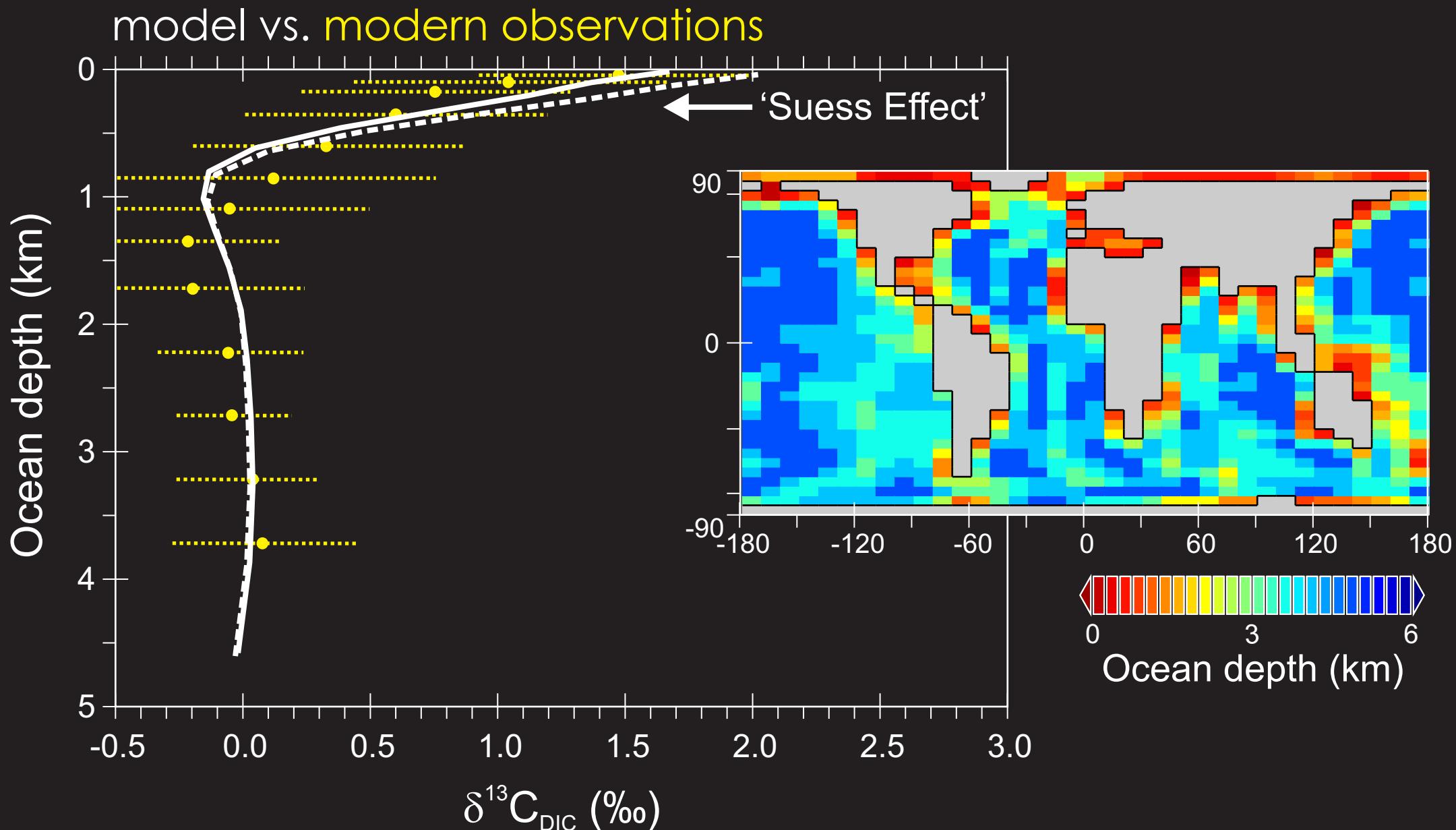
(biogenic mineral flux == 0)  
&&  
(organic carbon flux == 0)  
=> ballasting == .true.



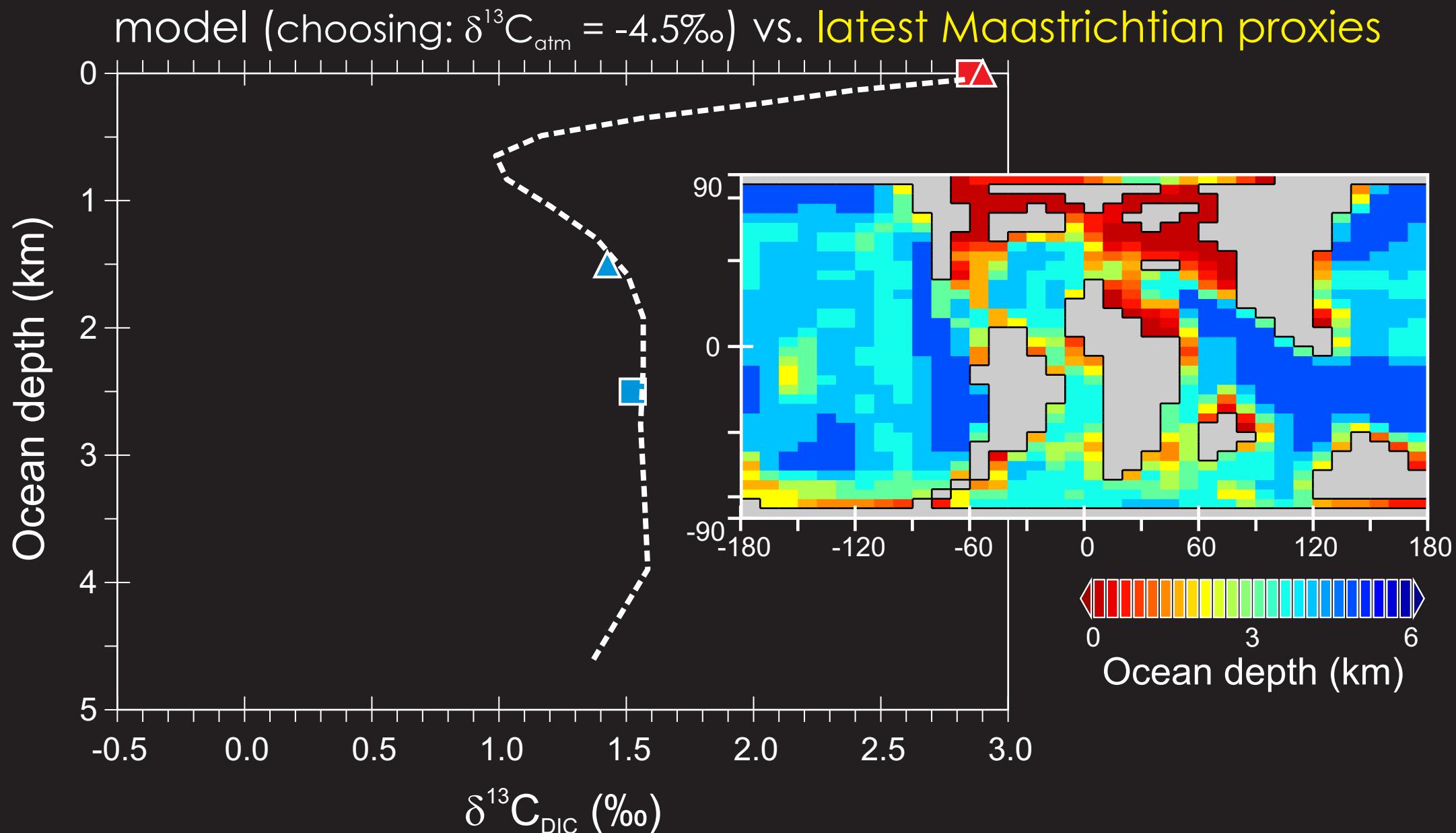
# *Evolution of the Biological Pump: ‘Hiccups’*



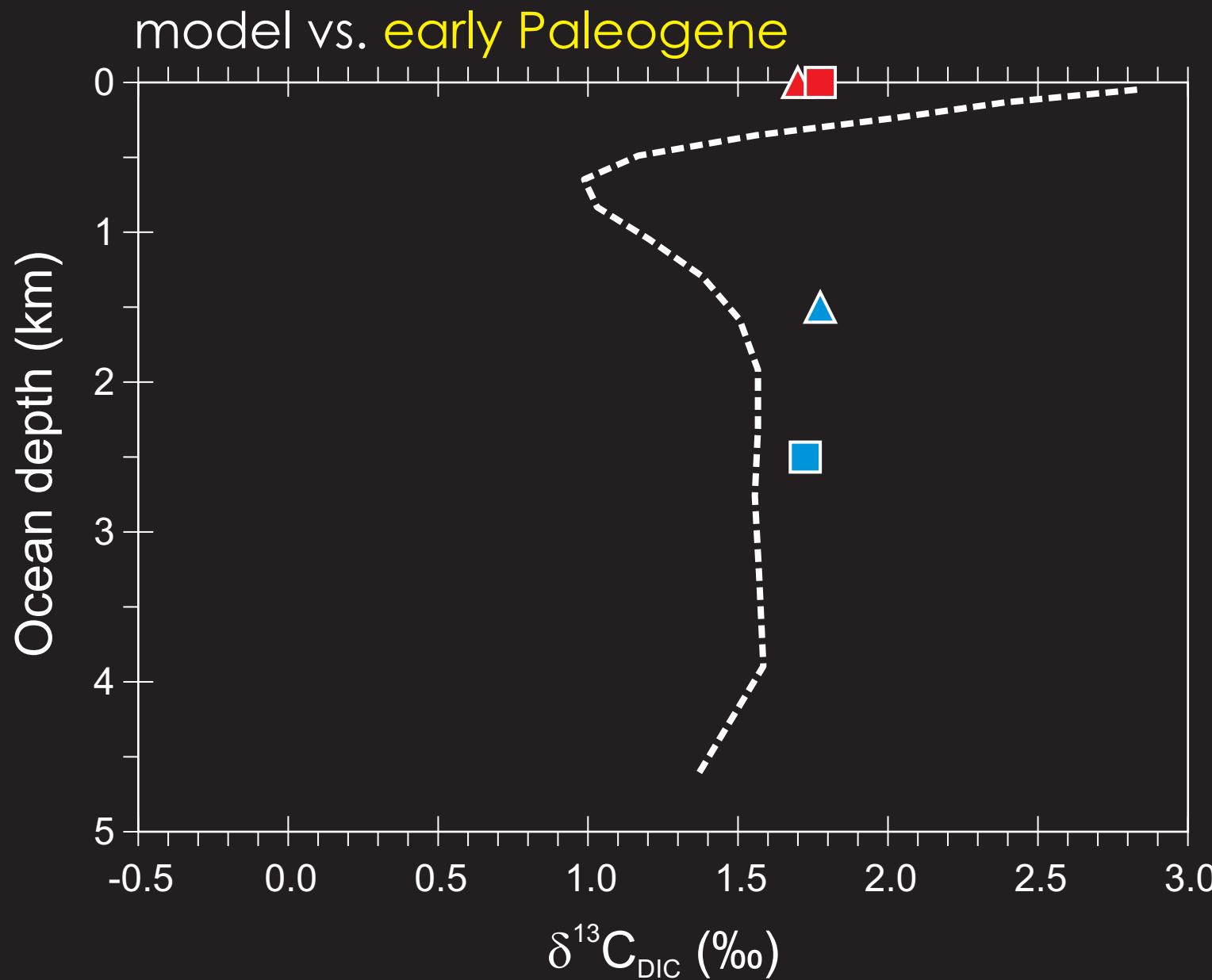
# *Evolution of the Biological Pump: ‘Hiccups’*



# *Evolution of the Biological Pump: ‘Hiccups’*



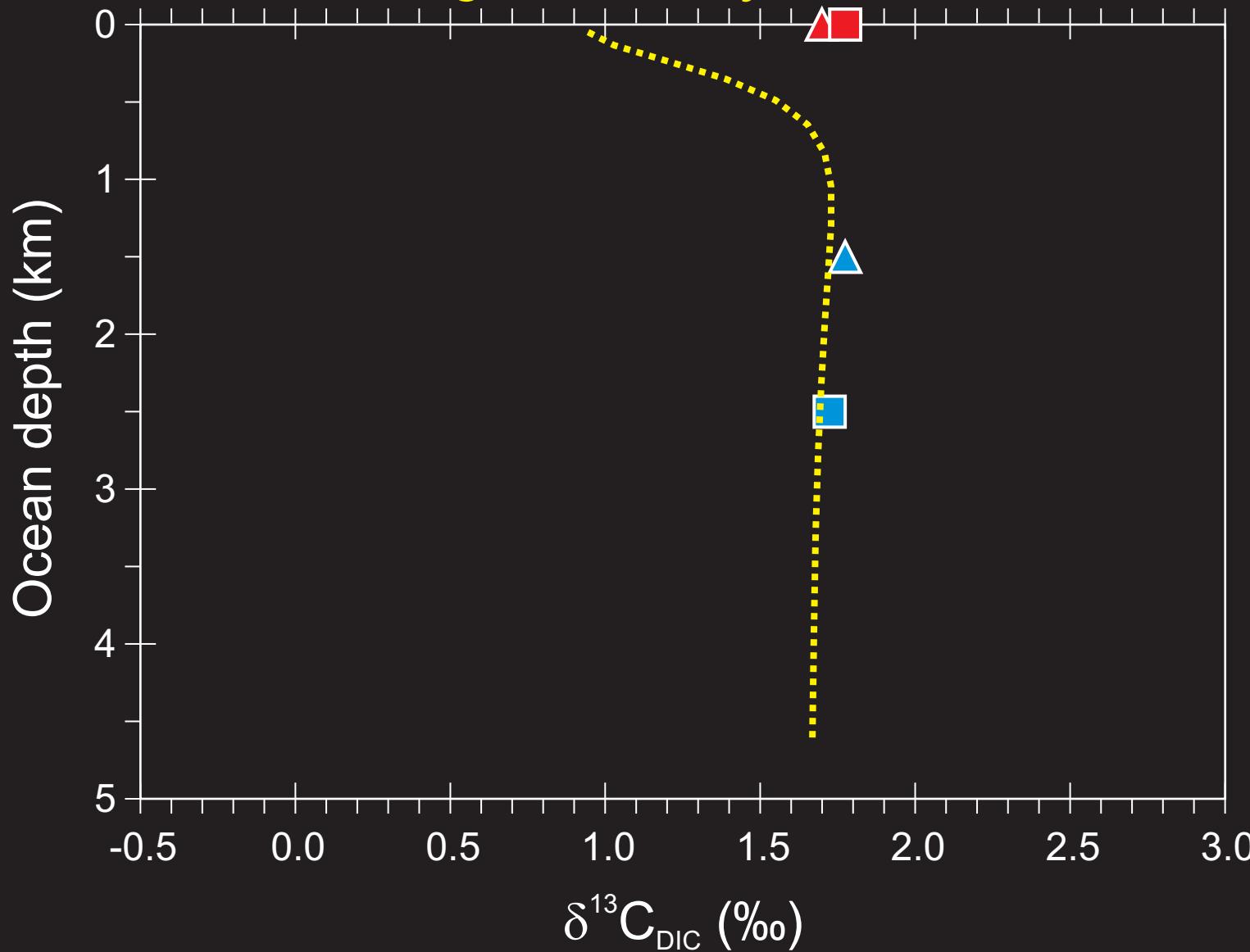
# *Evolution of the Biological Pump: ‘Hiccups’*



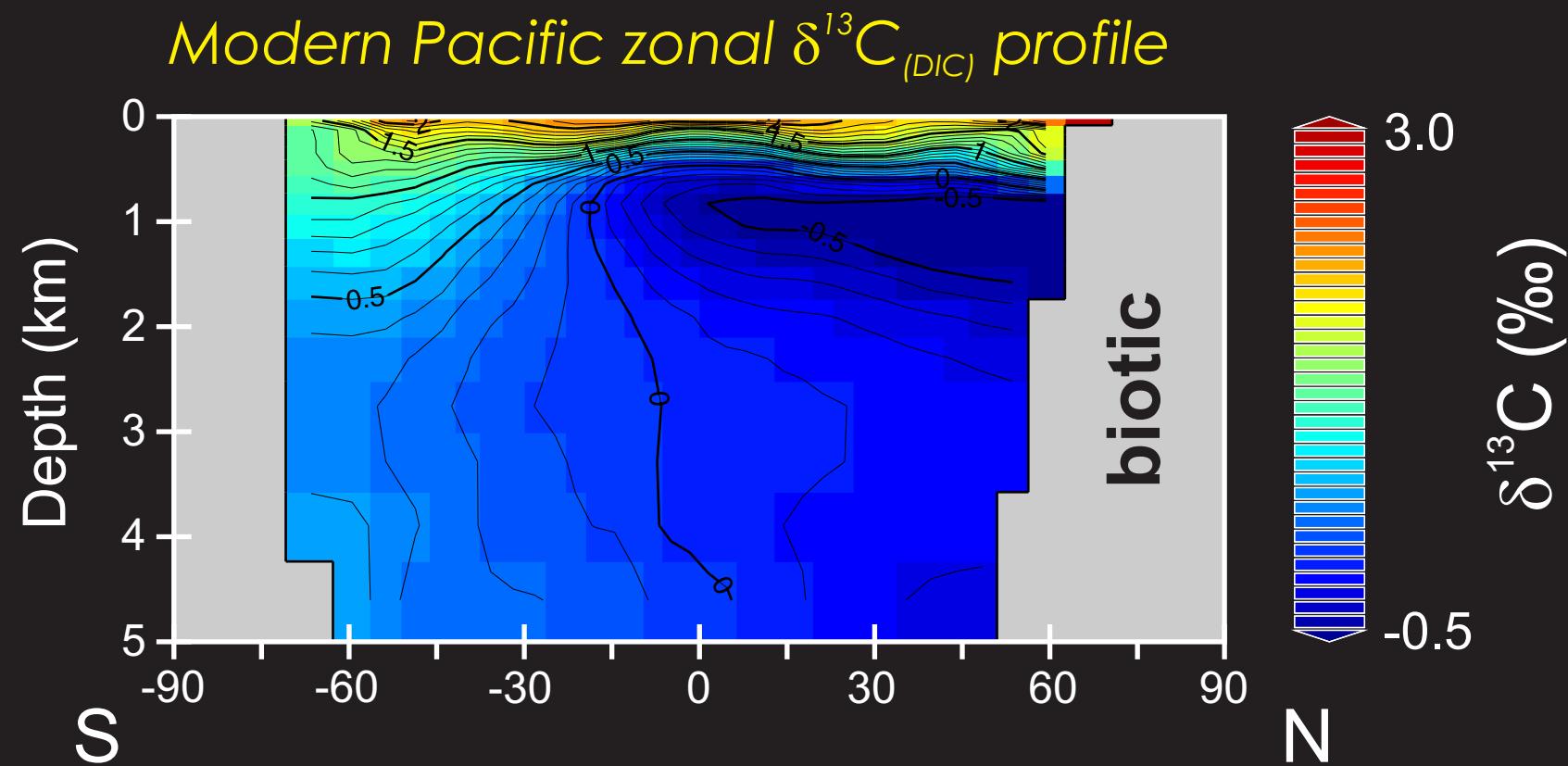
# *Evolution of the Biological Pump: ‘Hiccups’*



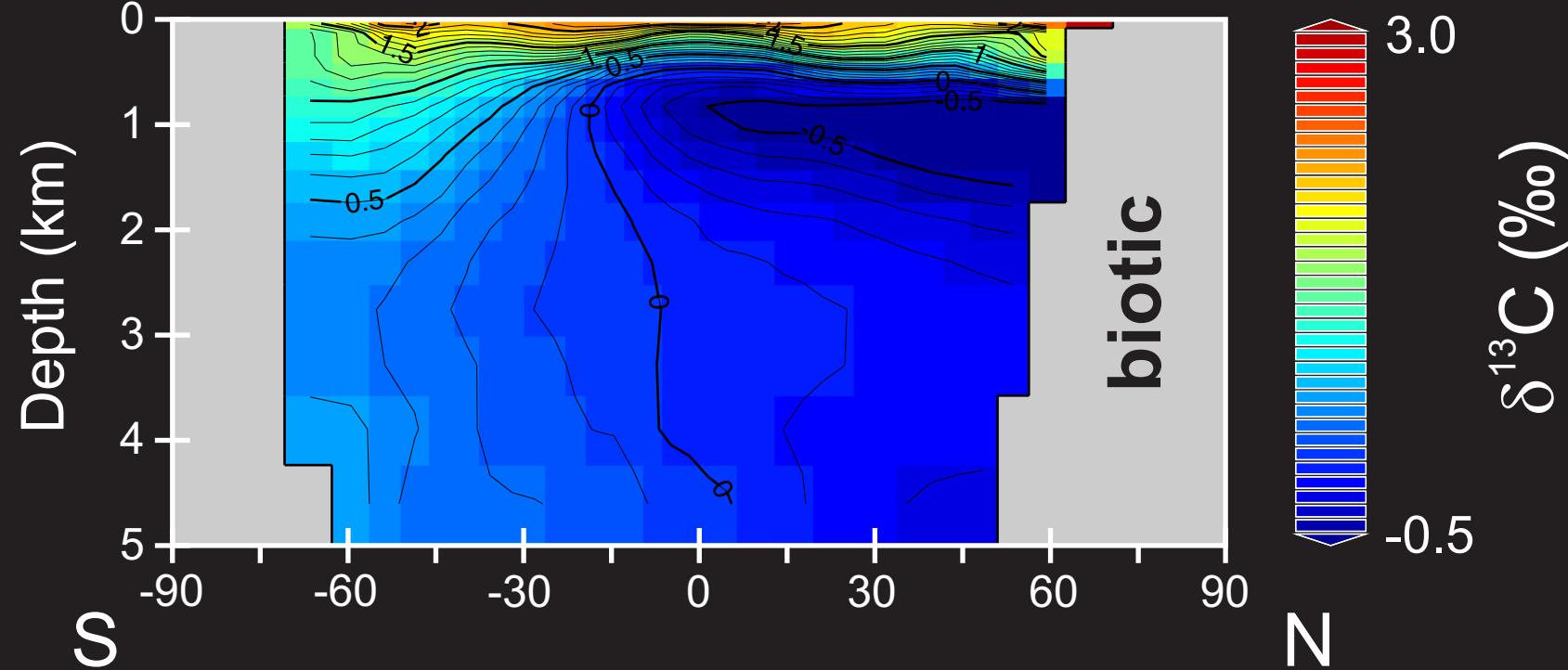
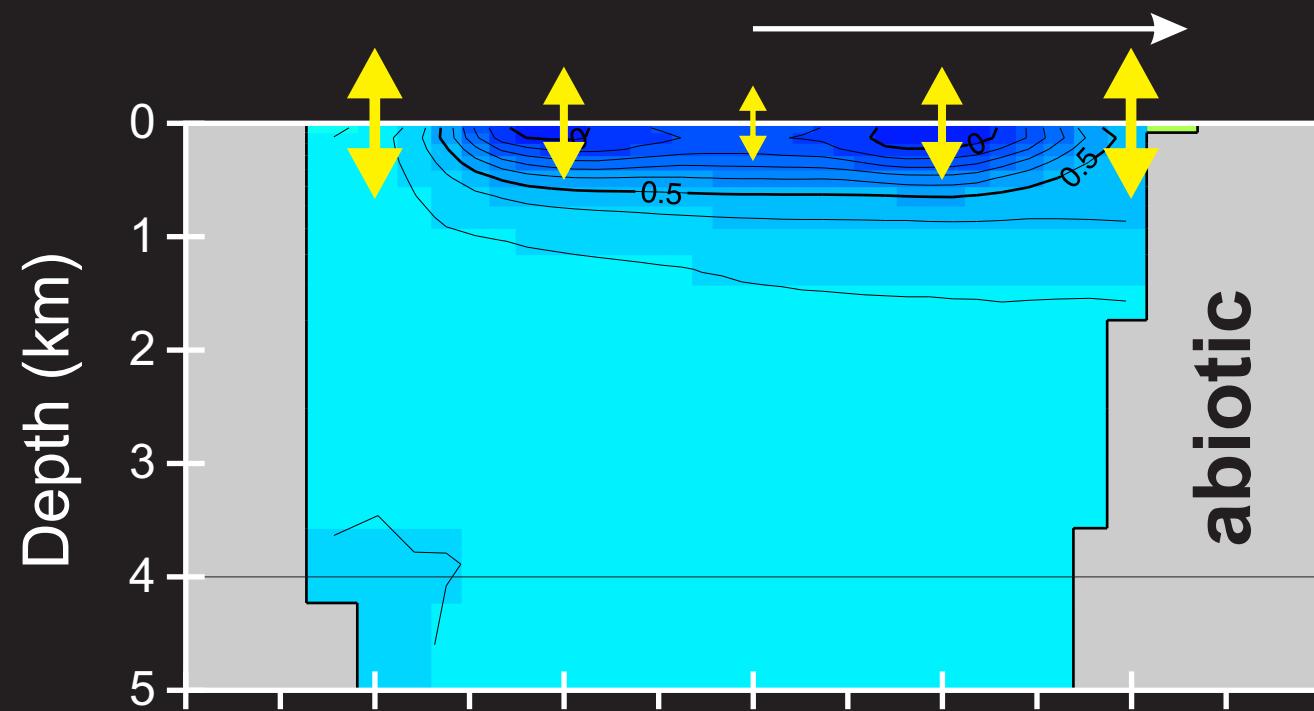
*0% biological activity*



# *Evolution of the Biological Pump: ‘Hiccups’*

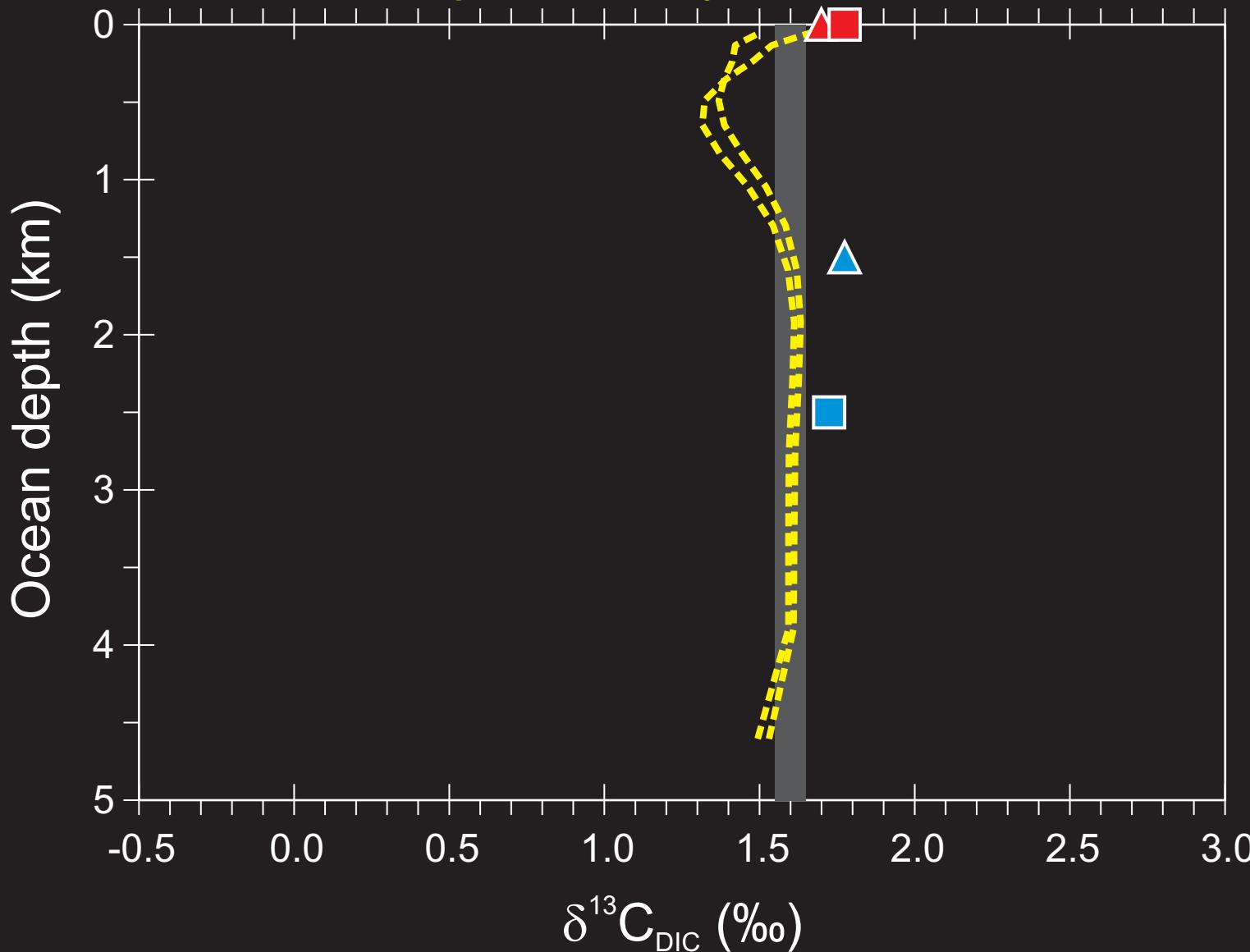


increasing fractionation between  $p\text{CO}_2$  and  $[\text{CO}_2]$   
with decreasing temperature towards to poles



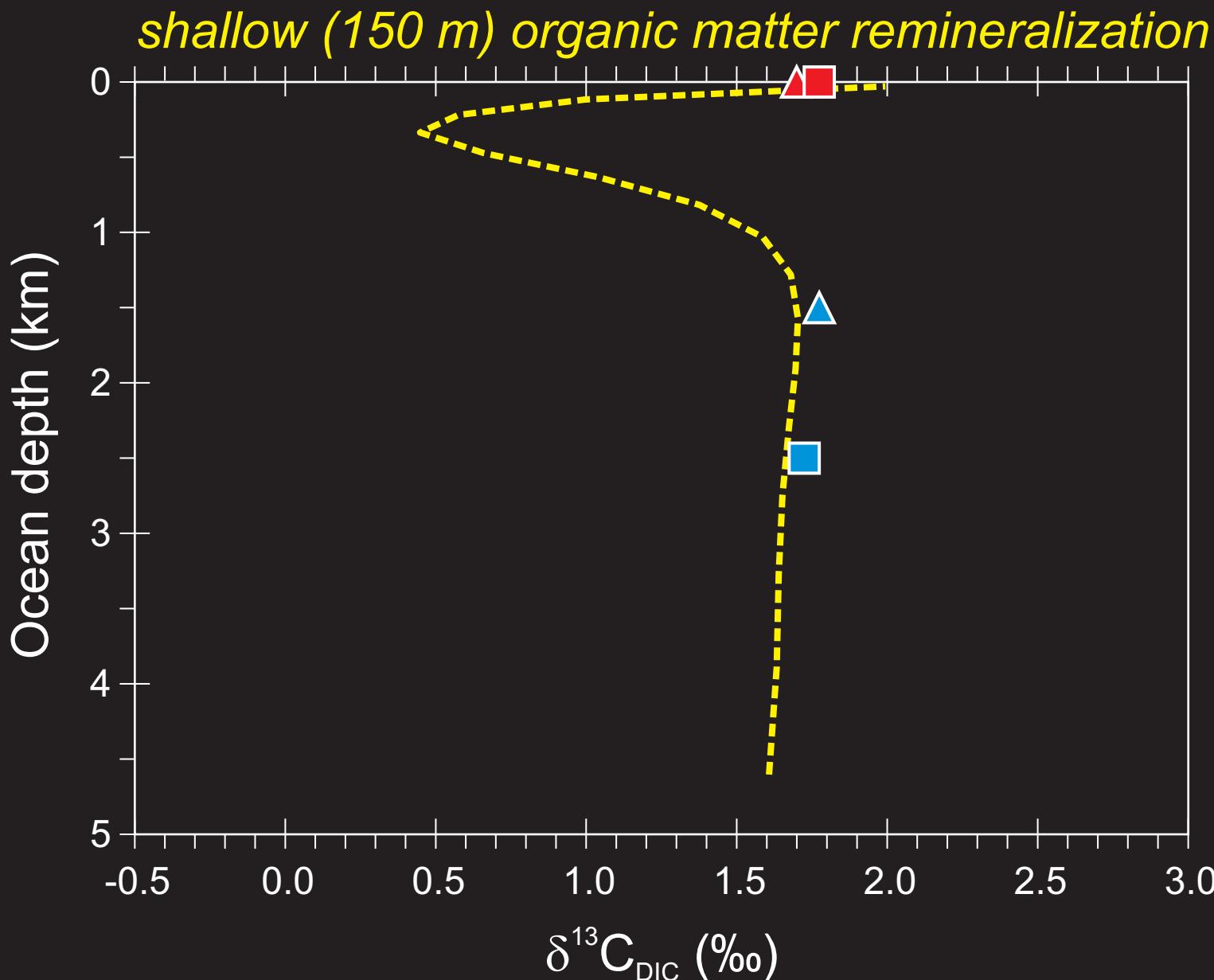
# *Evolution of the Biological Pump: ‘Hiccups’*

*30-40% biological activity*



Answer:  
A somewhat  
reduced biological  
pump ...

# *Evolution of the Biological Pump: ‘Hiccups’*

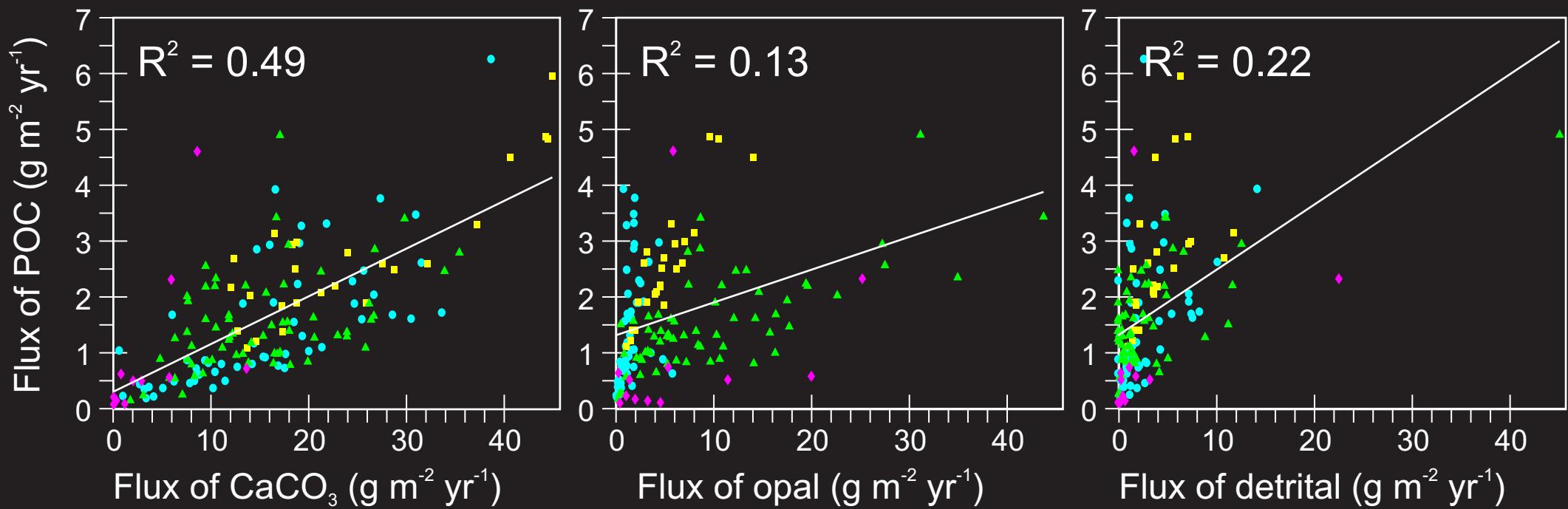


Answer:  
A somewhat  
reduced biological  
pump ...  
... or, a strange and  
different biological  
pump, consistent  
with profound  
ecological change  
post impact?

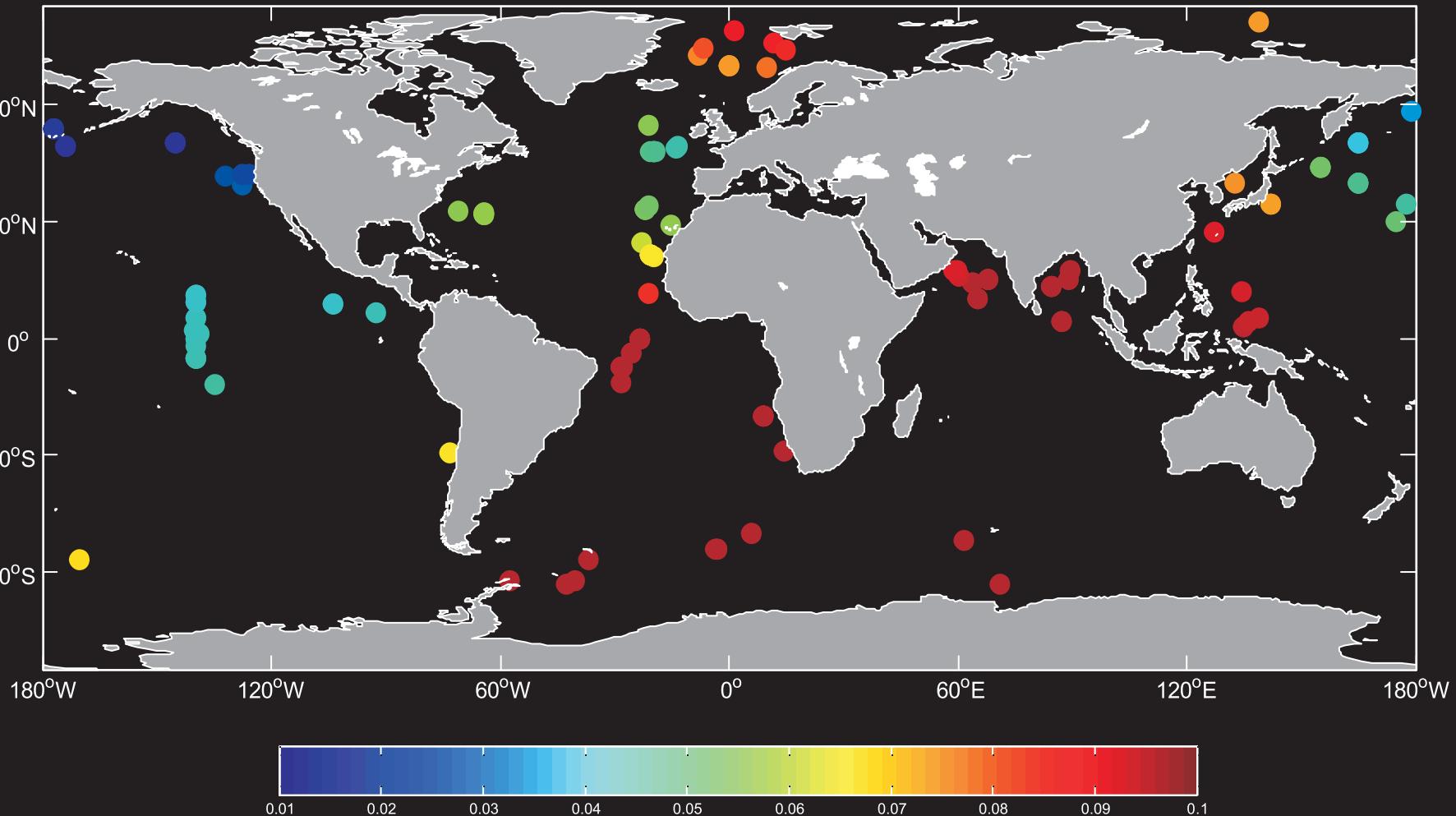
# *Evolution of the Biological Pump:* Planktic carbonate production and ‘ballasting’

Compilation of sediment trap observations:  
depths  $\geq 2000$  m to exclude hydrodynamically distorted  
fluxes and relationships, and differentiated by basin:  
cyan == Atlantic, yellow == Indian, green == Pacific,  
magenta == Southern Ocean.

[Wilson et al., 2012; GBC 26, doi:10.1029/2012GB004398]



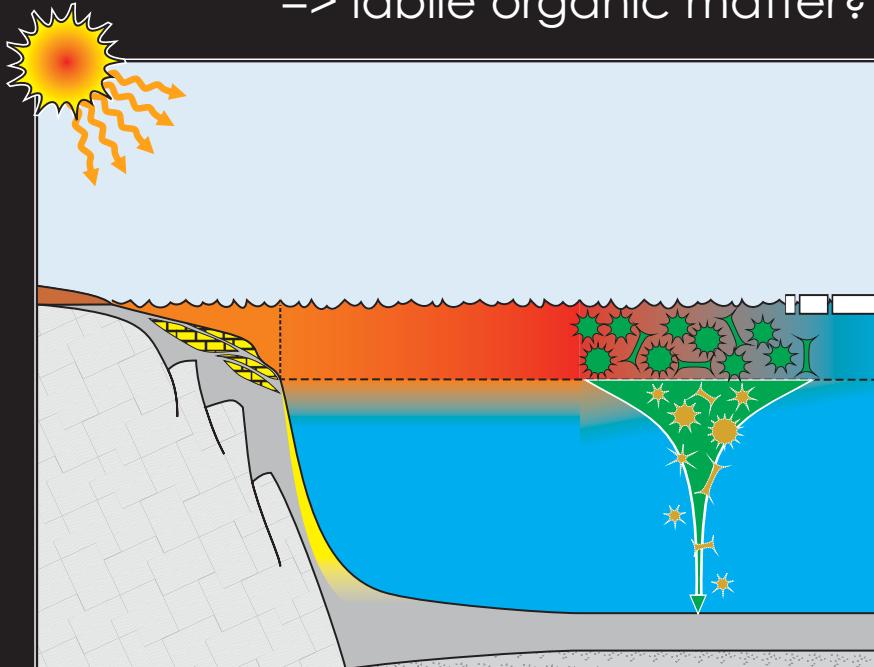
# *Evolution of the Biological Pump:* Planktic carbonate production and ‘ballasting’



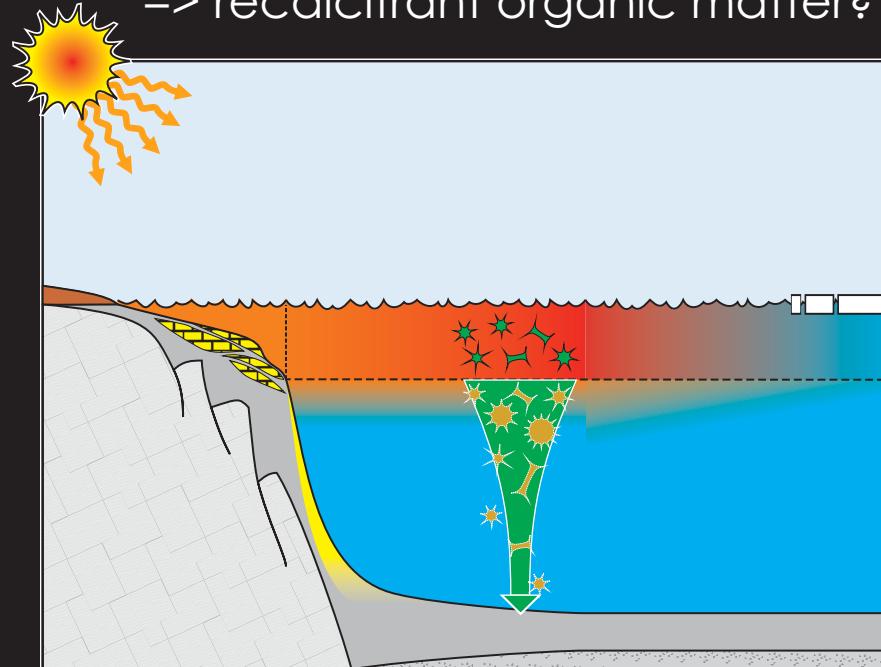
Spatial distribution of carrying capacity (ballasting) coefficients  
calculated using geographically weighted regression  
analysis for  $\text{CaCO}_3$ .

# *Evolution of the Biological Pump:* Planktic carbonate production and ‘ballasting’

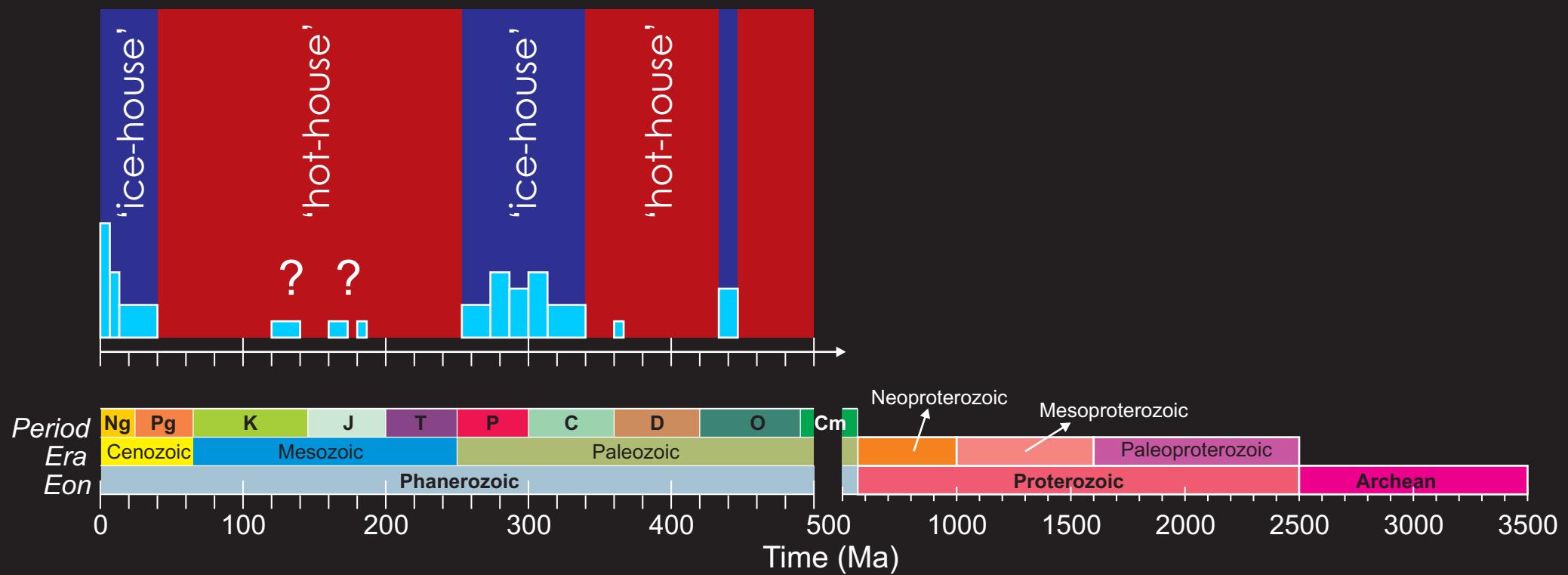
High productivity and export efficiency, but low biological pump efficiency.  
=> labile organic matter?



Low productivity and export efficiency, but high biological pump efficiency.  
=> recalcitrant organic matter?



# *Evolution of the Biological Pump:* Ocean Carbon Cycling and Oxygenation in Warm Climates

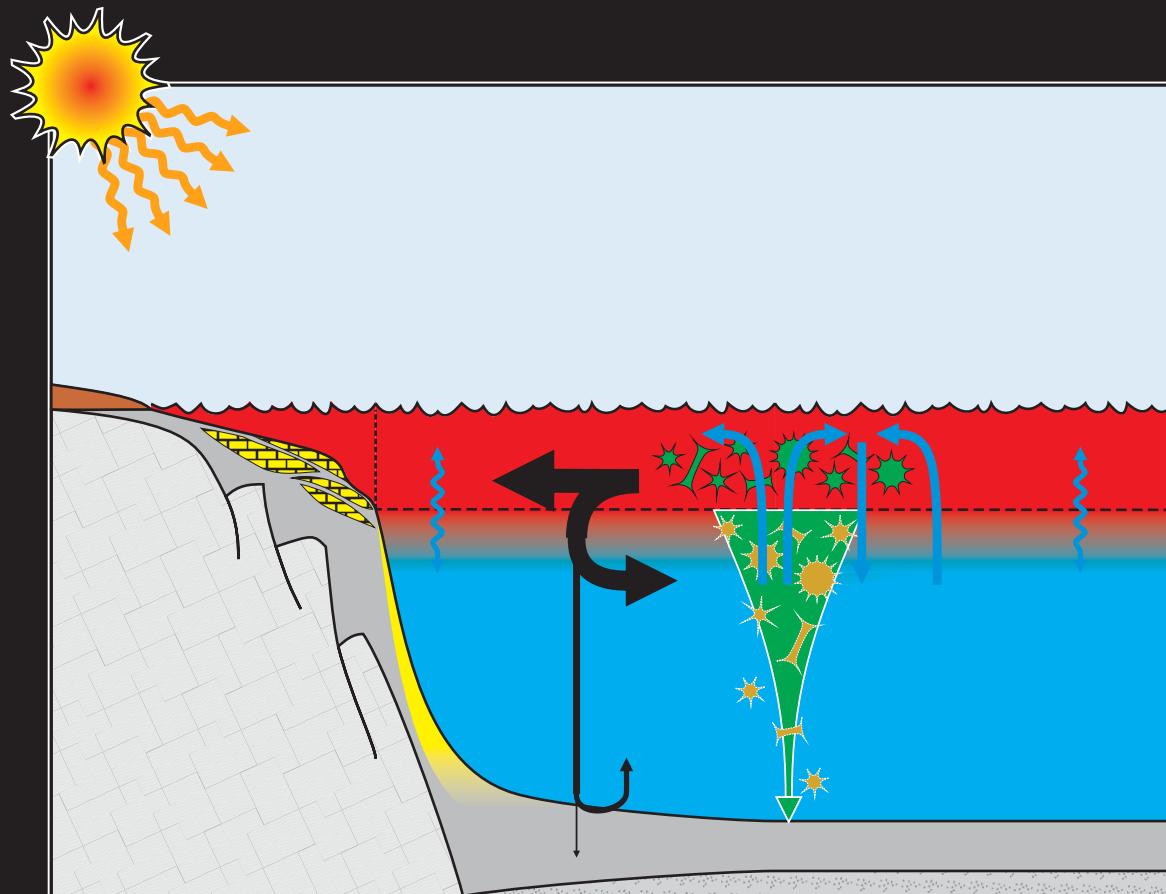


# *Evolution of the Biological Pump:* Ocean Carbon Cycling and Oxygenation in Warm Climates

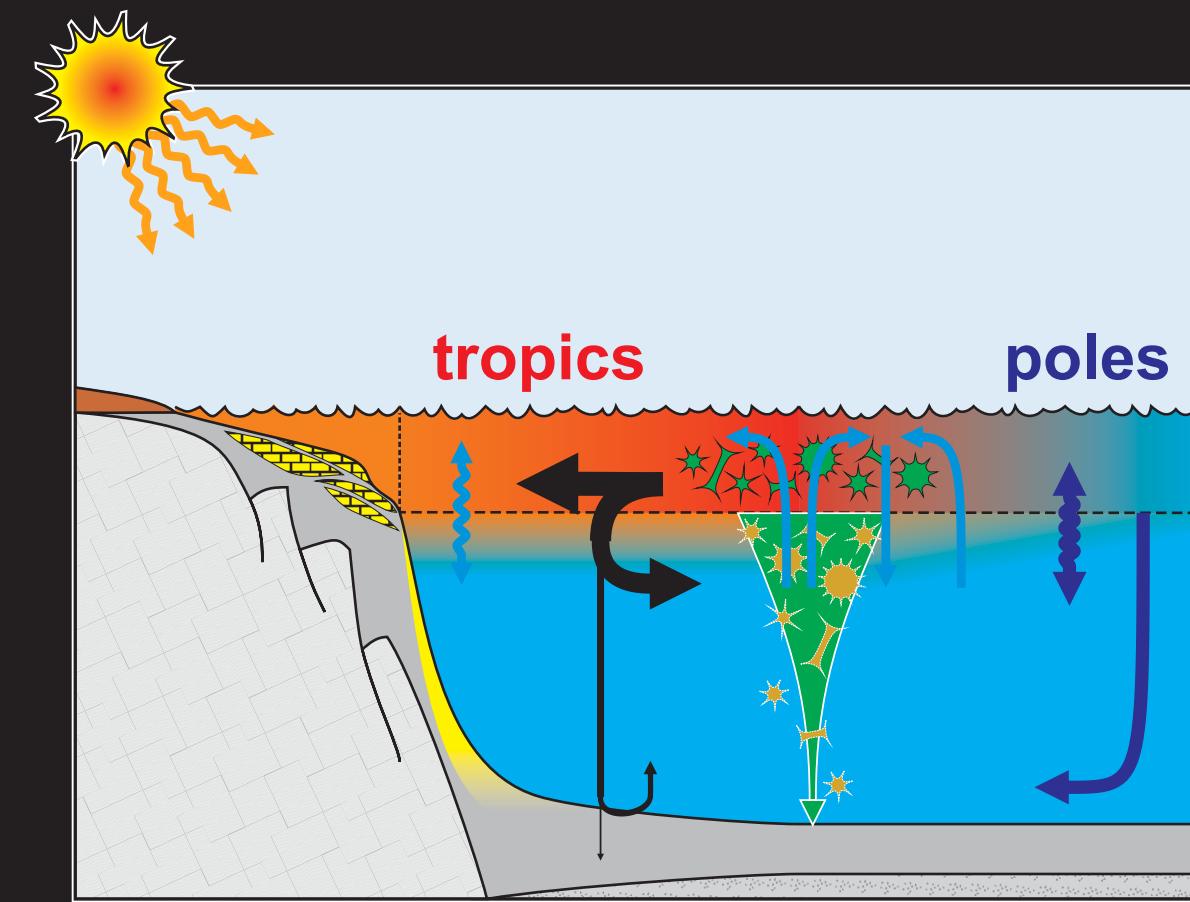
(warm == stratified) && (stratified == anoxic) == .true.

???

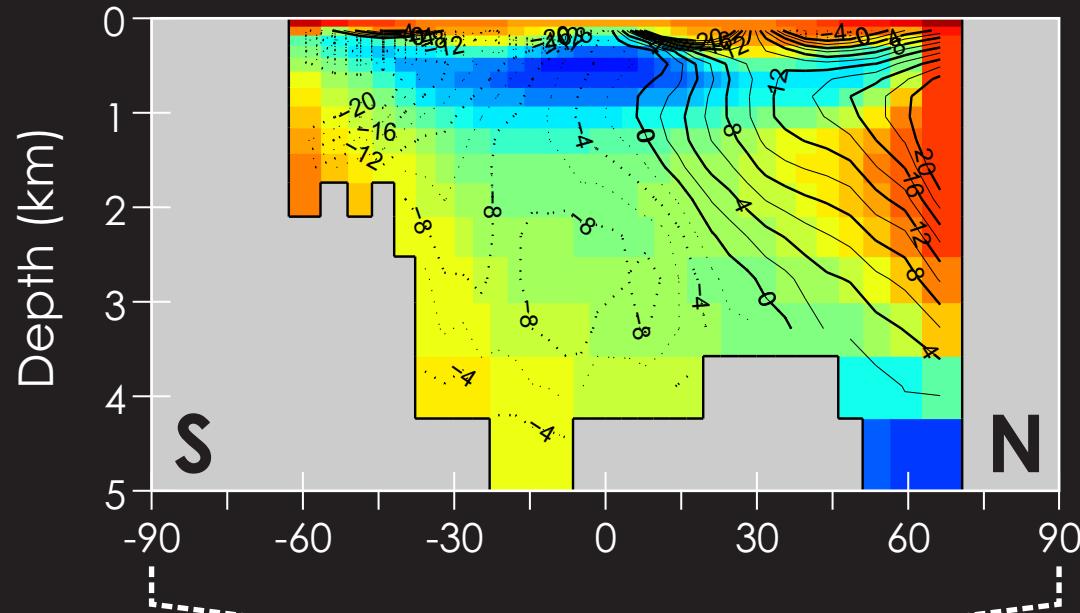
('stratified' || 'sluggish' || 'stagnant')



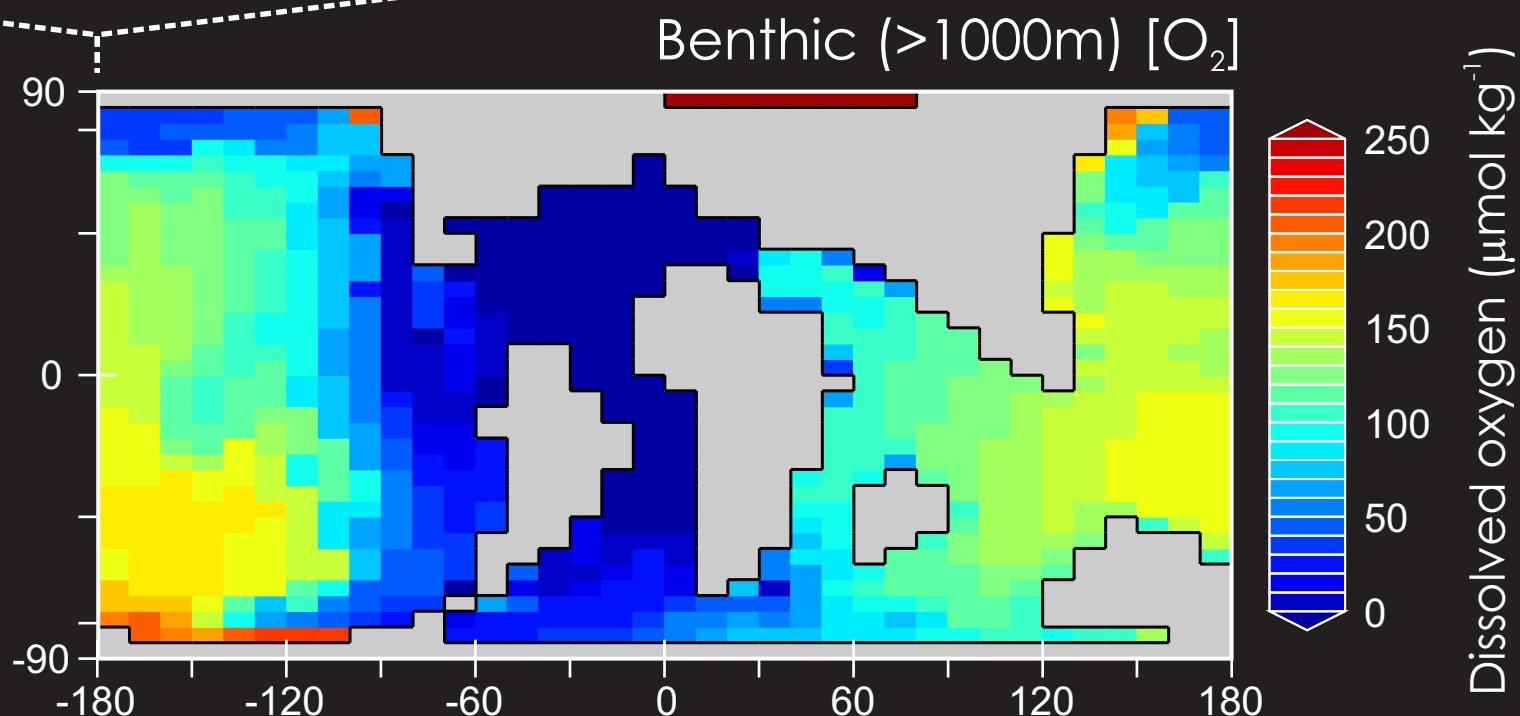
# *Evolution of the Biological Pump:* Ocean Carbon Cycling and Oxygenation in Warm Climates



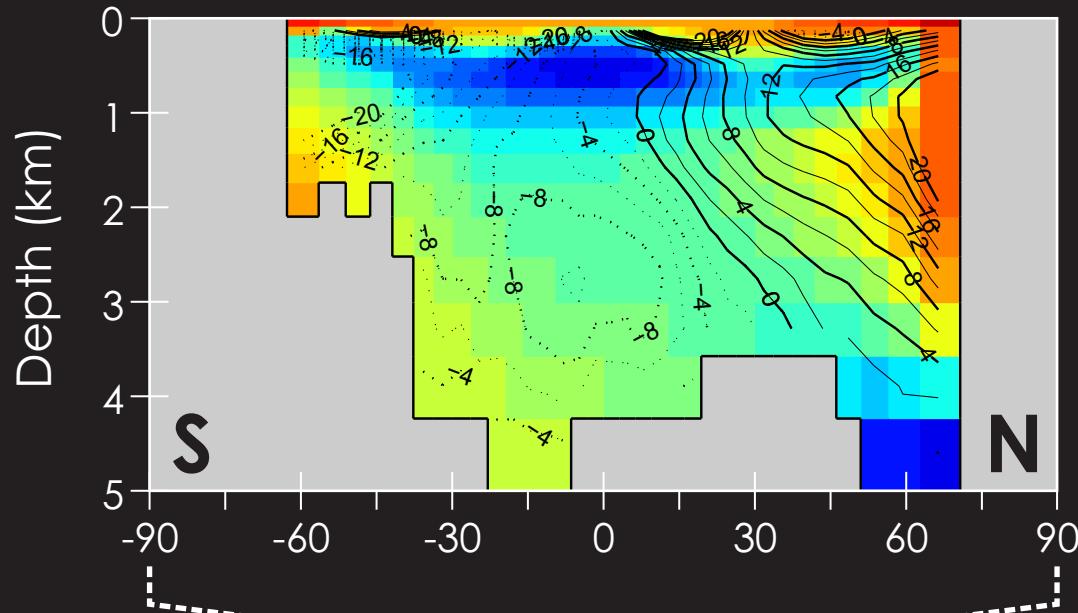
# *Evolution of the Biological Pump: Ocean Carbon Cycling and Oxygenation in Warm Climates*



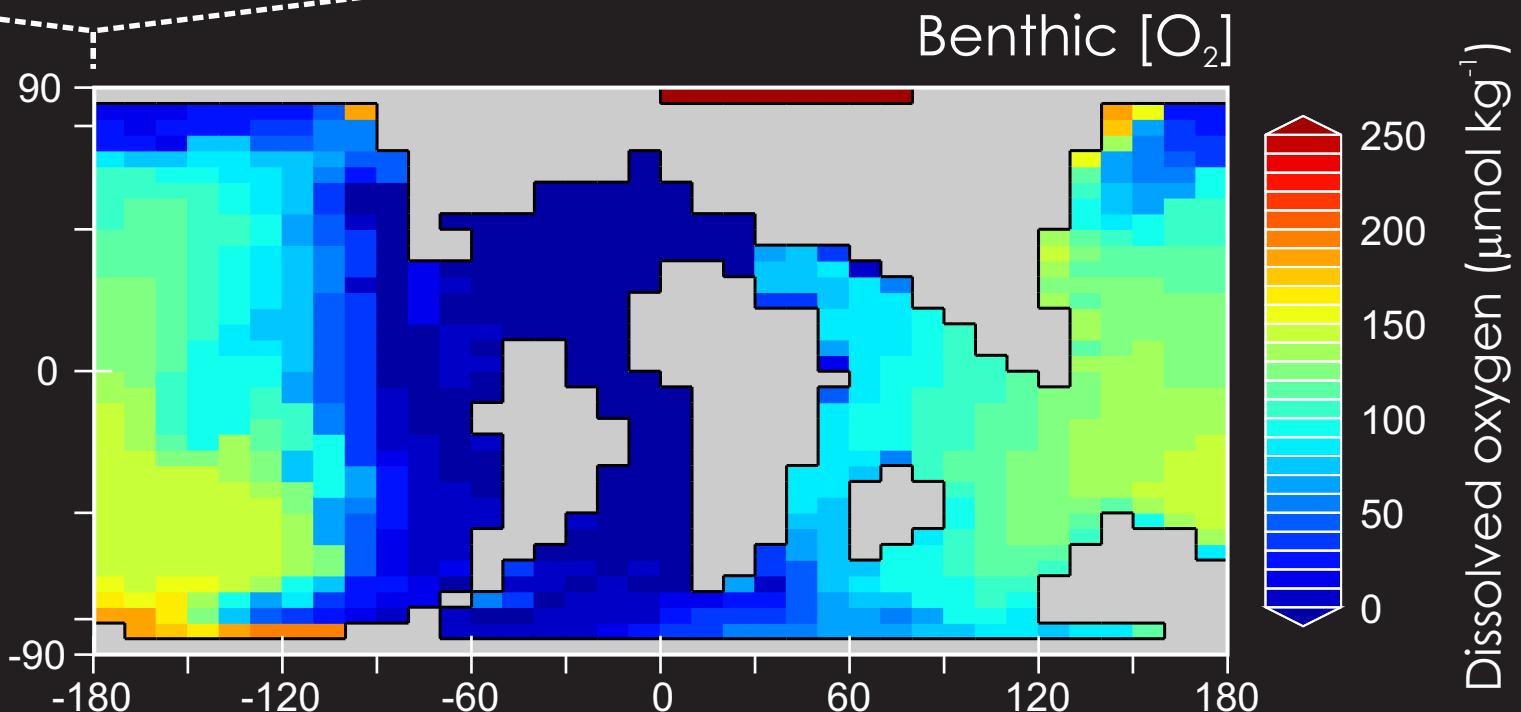
x4 CO<sub>2</sub> reference simulation



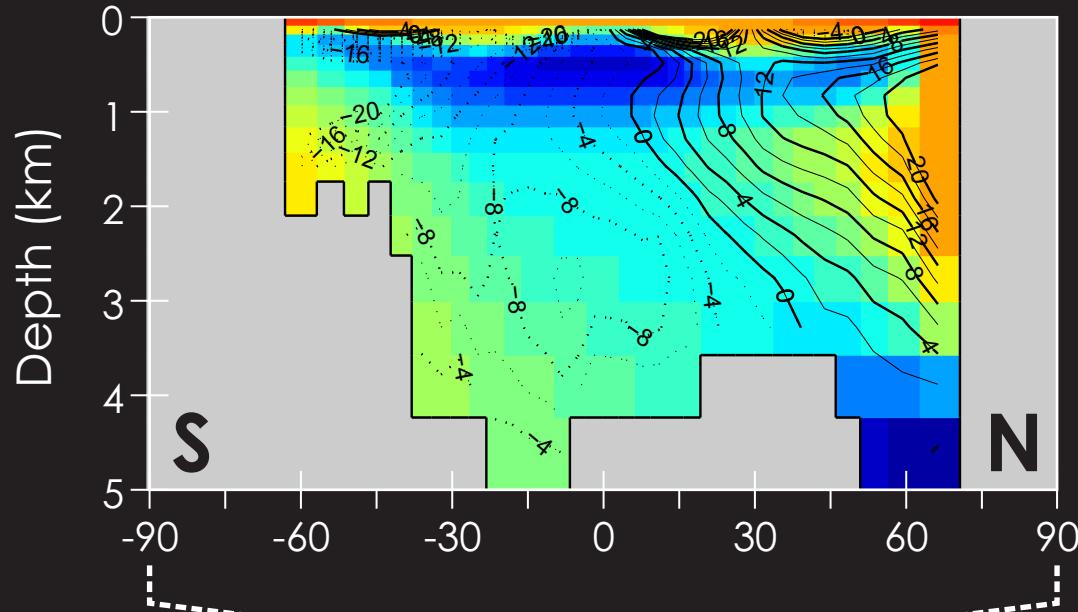
# *Evolution of the Biological Pump:* Ocean Carbon Cycling and Oxygenation in Warm Climates



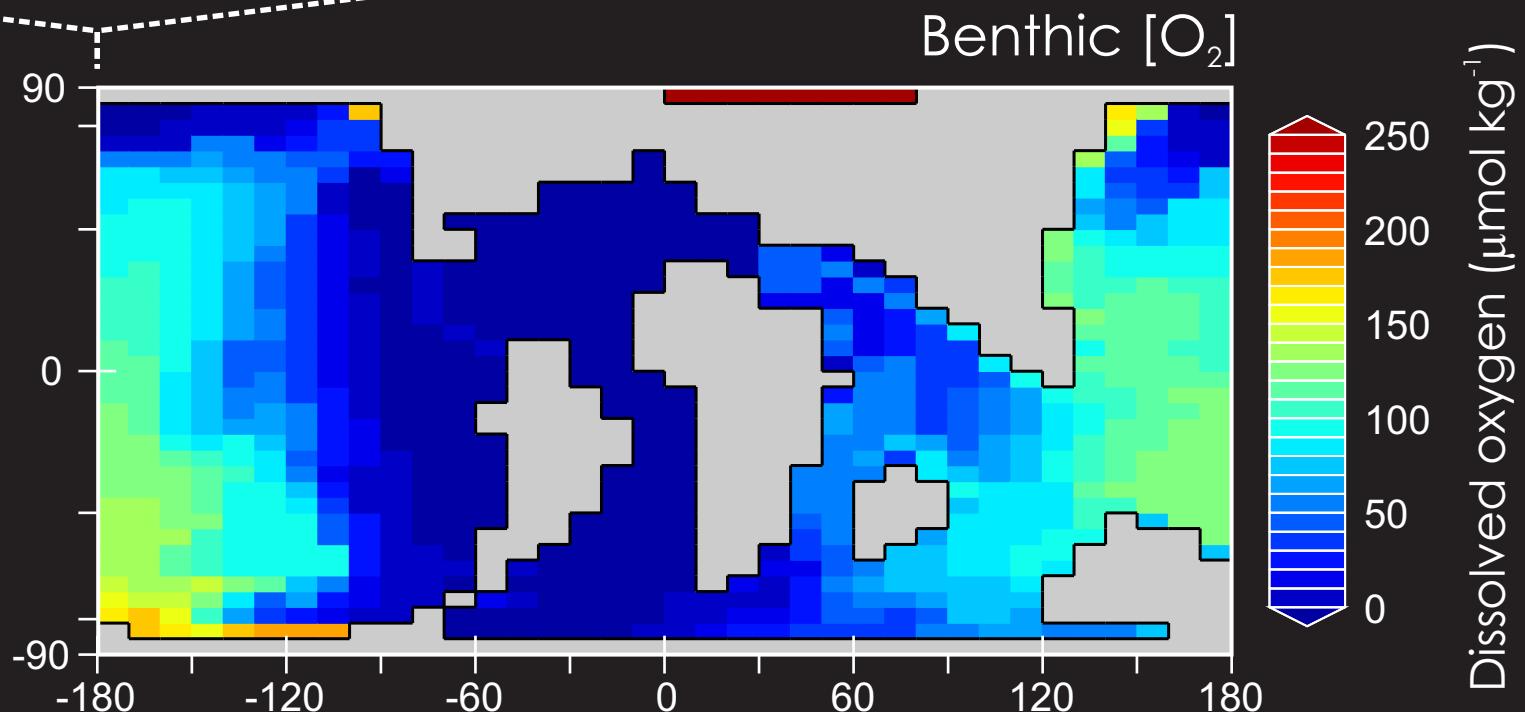
x8 CO<sub>2</sub> @ 10,000 yrs  
(started from end of the x4 simulation)



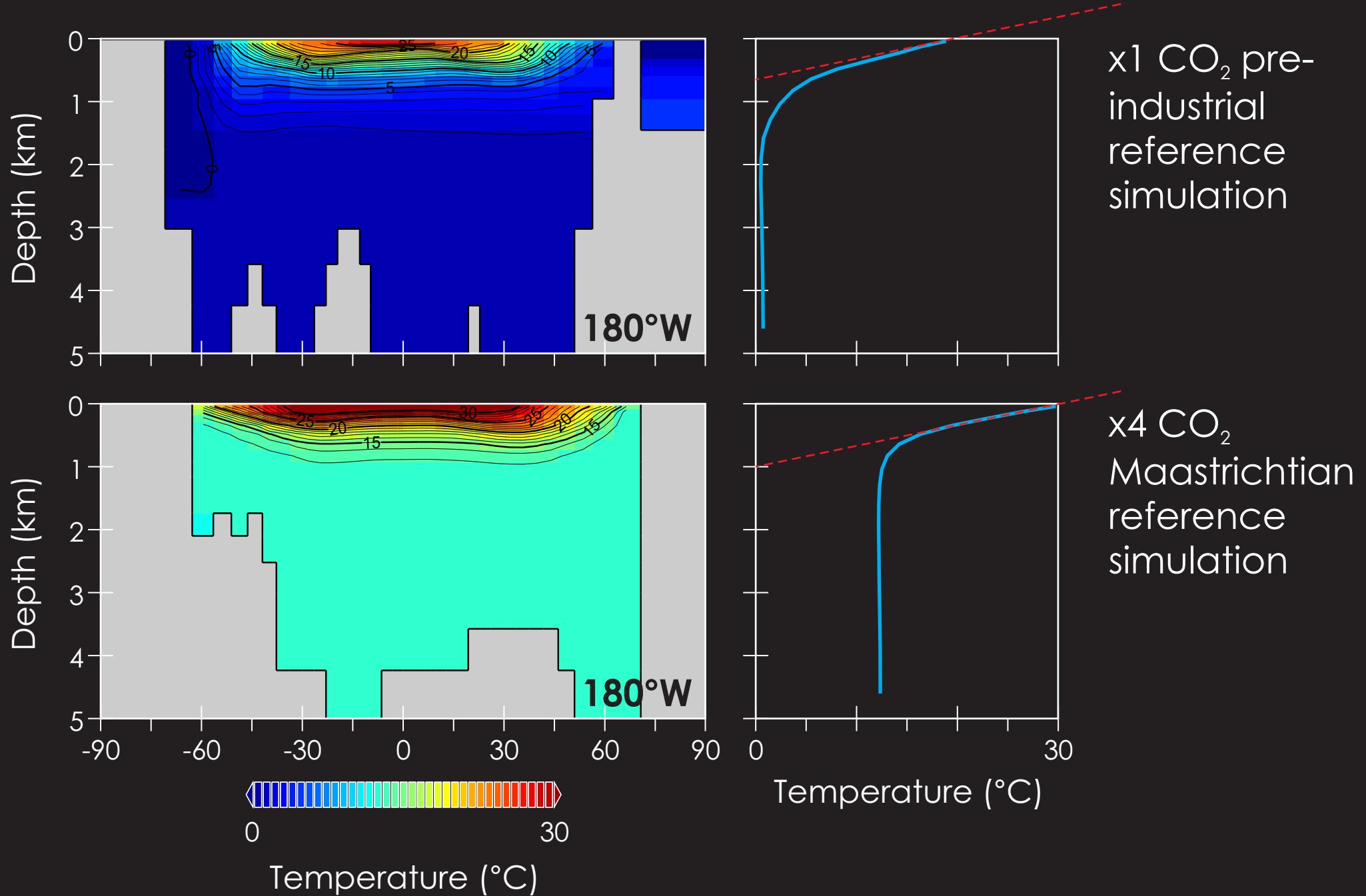
# *Evolution of the Biological Pump:* Ocean Carbon Cycling and Oxygenation in Warm Climates



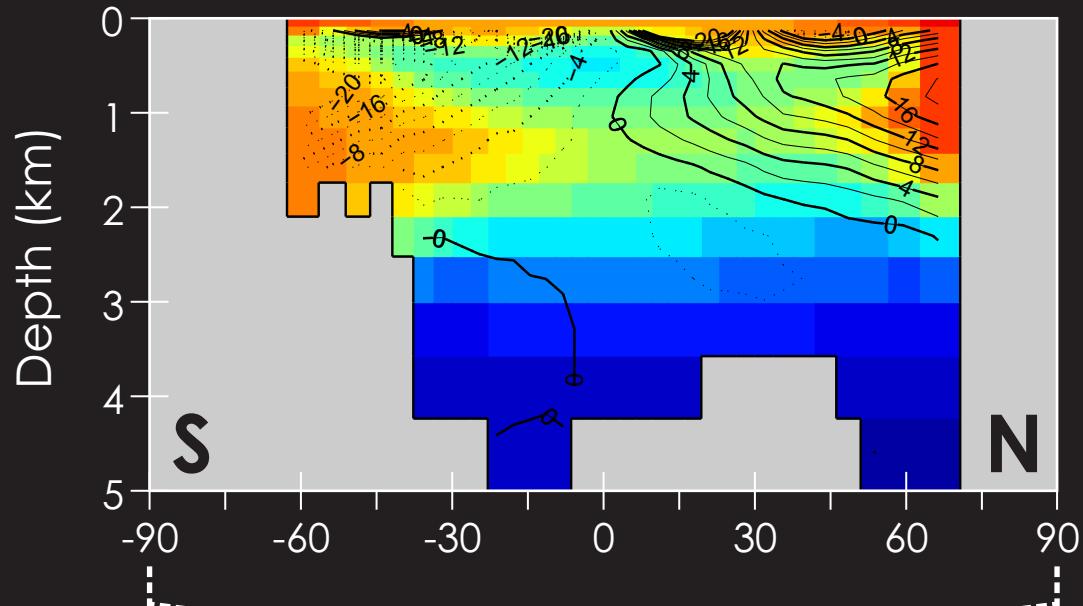
$\times 16 \text{ CO}_2 @ 10,000 \text{ yrs}$   
(started from end of the x4 simulation)



# *Evolution of the Biological Pump:* Ocean Carbon Cycling and Oxygenation in Warm Climates

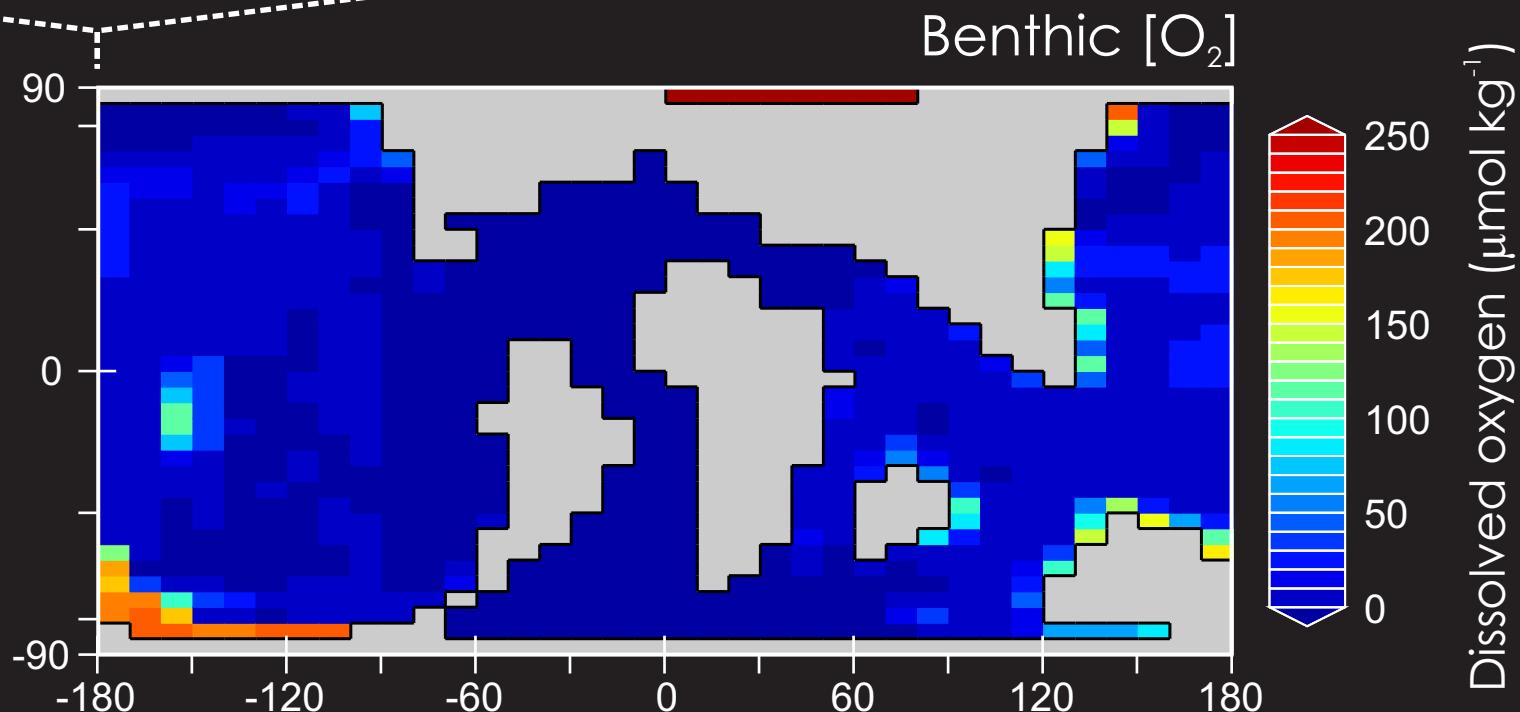


# *Evolution of the Biological Pump:* Ocean Carbon Cycling and Oxygenation in Warm Climates

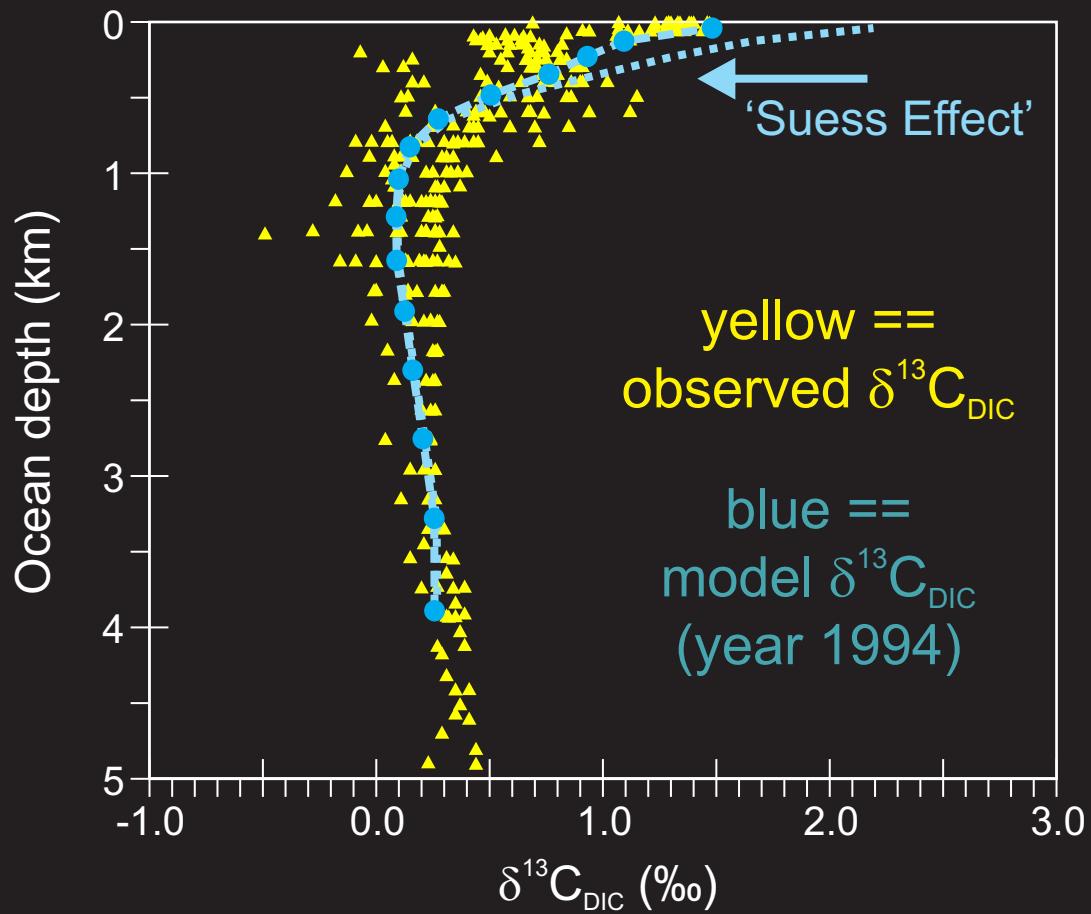
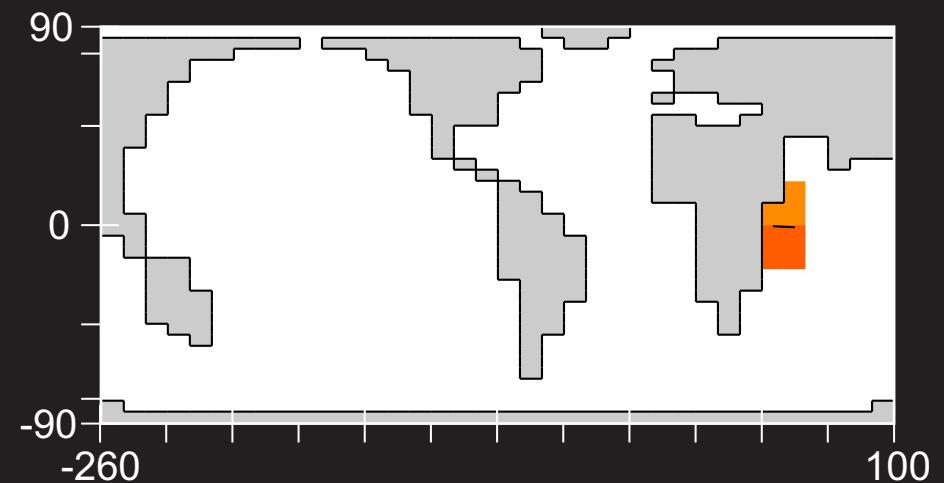


x16 CO<sub>2</sub> @ 2,000 yrs

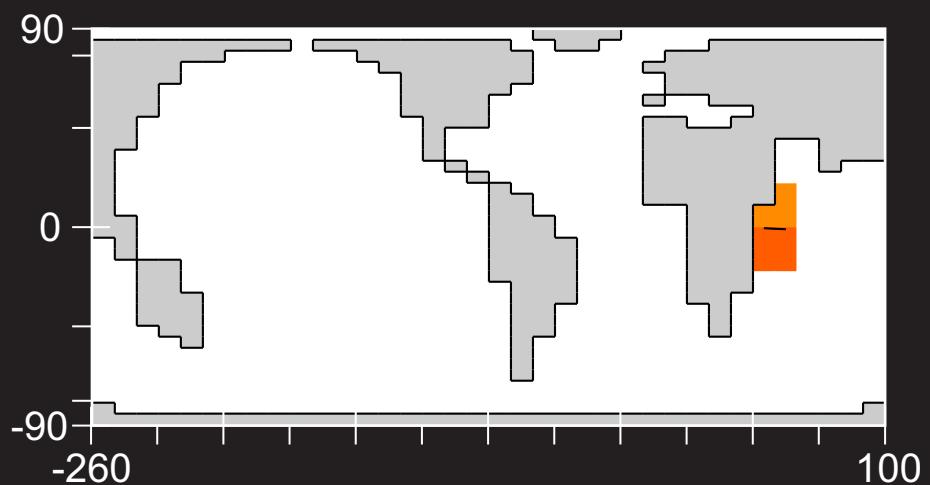
transient state  
(incomplete adjustment to  
increased radiative forcing)



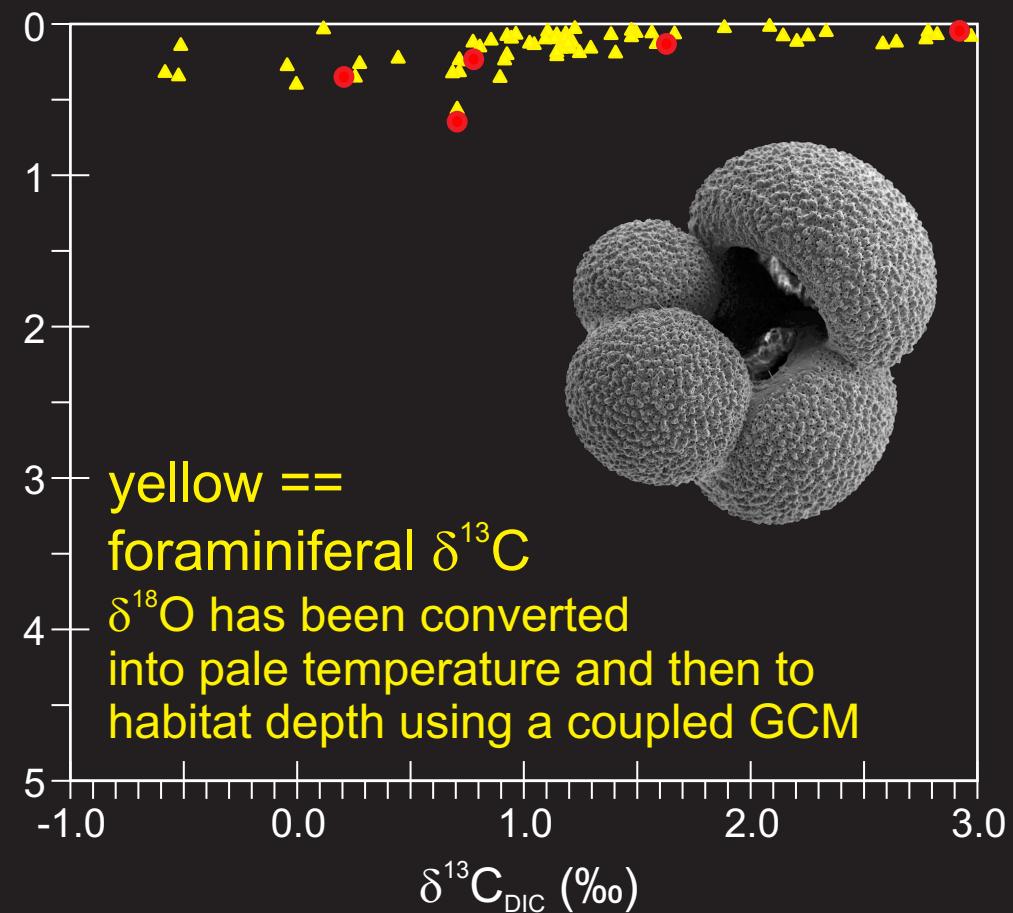
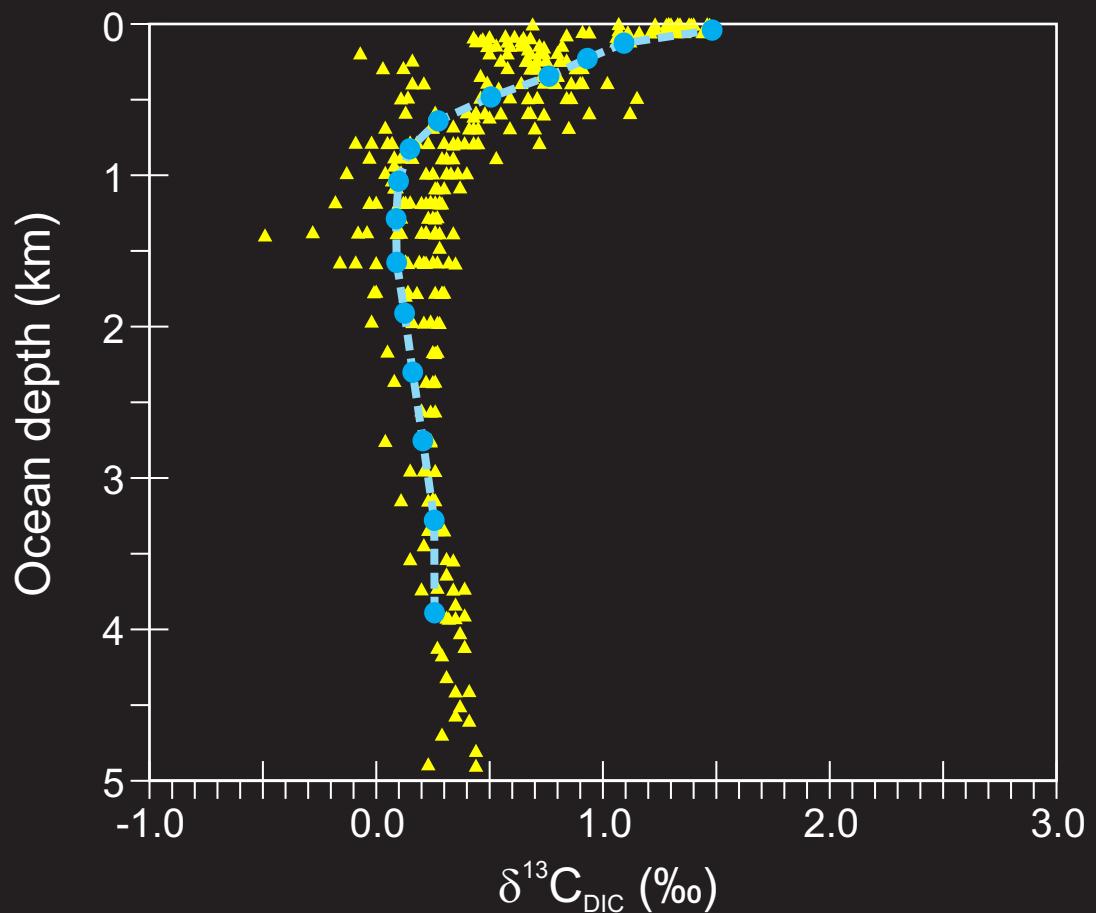
# Open ocean $\delta^{13}\text{C}_{\text{DIC}}$ adjacent to modern Tanzania



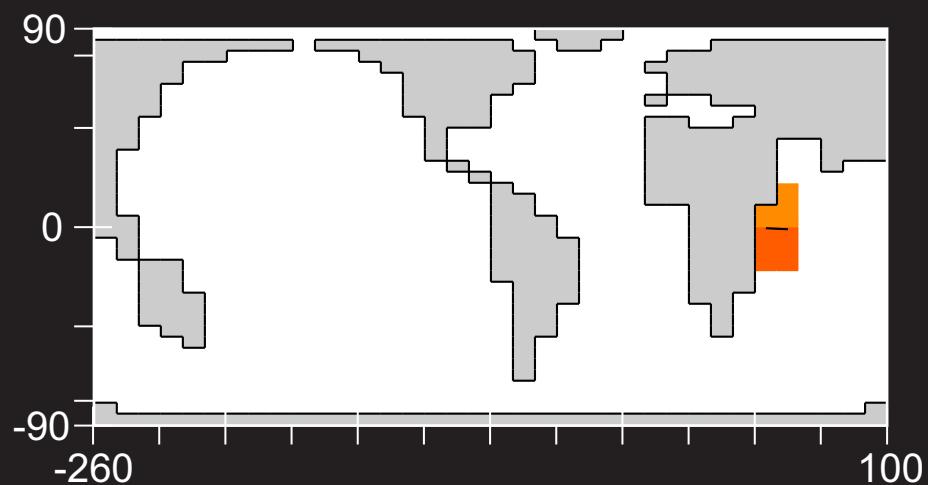
Open ocean  $\delta^{13}\text{C}_{\text{DIC}}$  adjacent to modern Tanzania



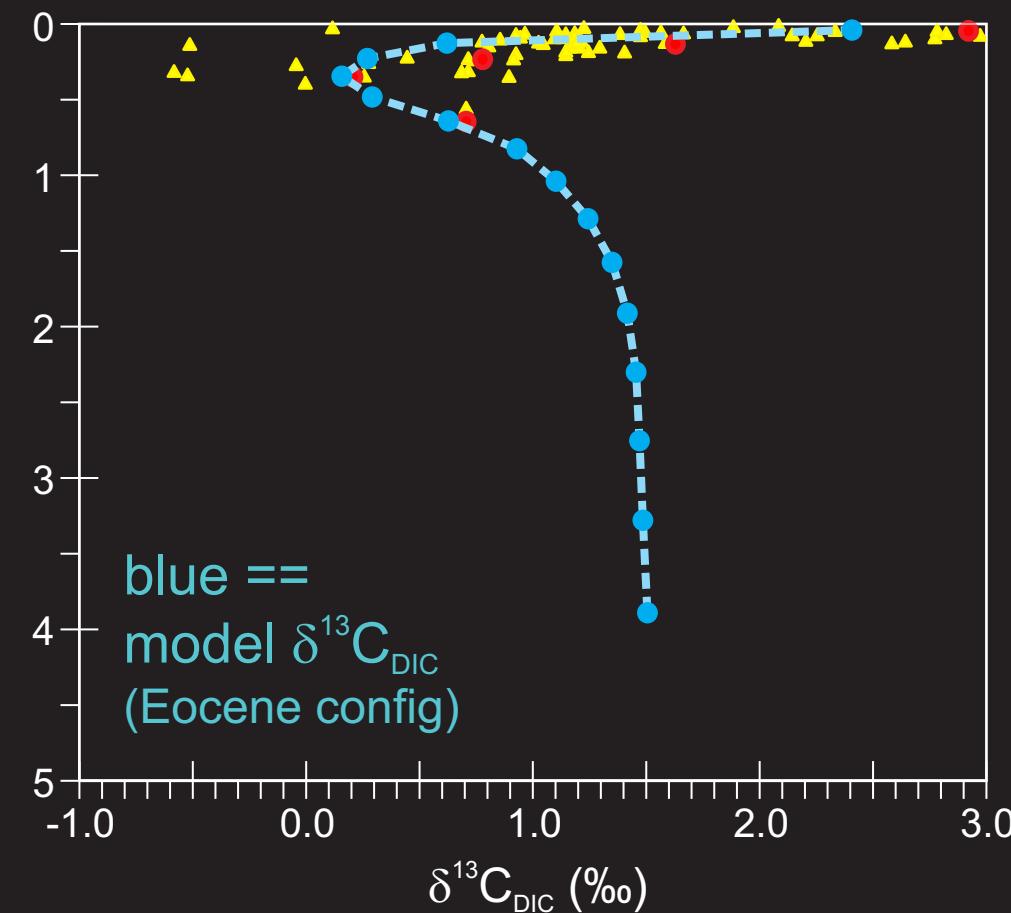
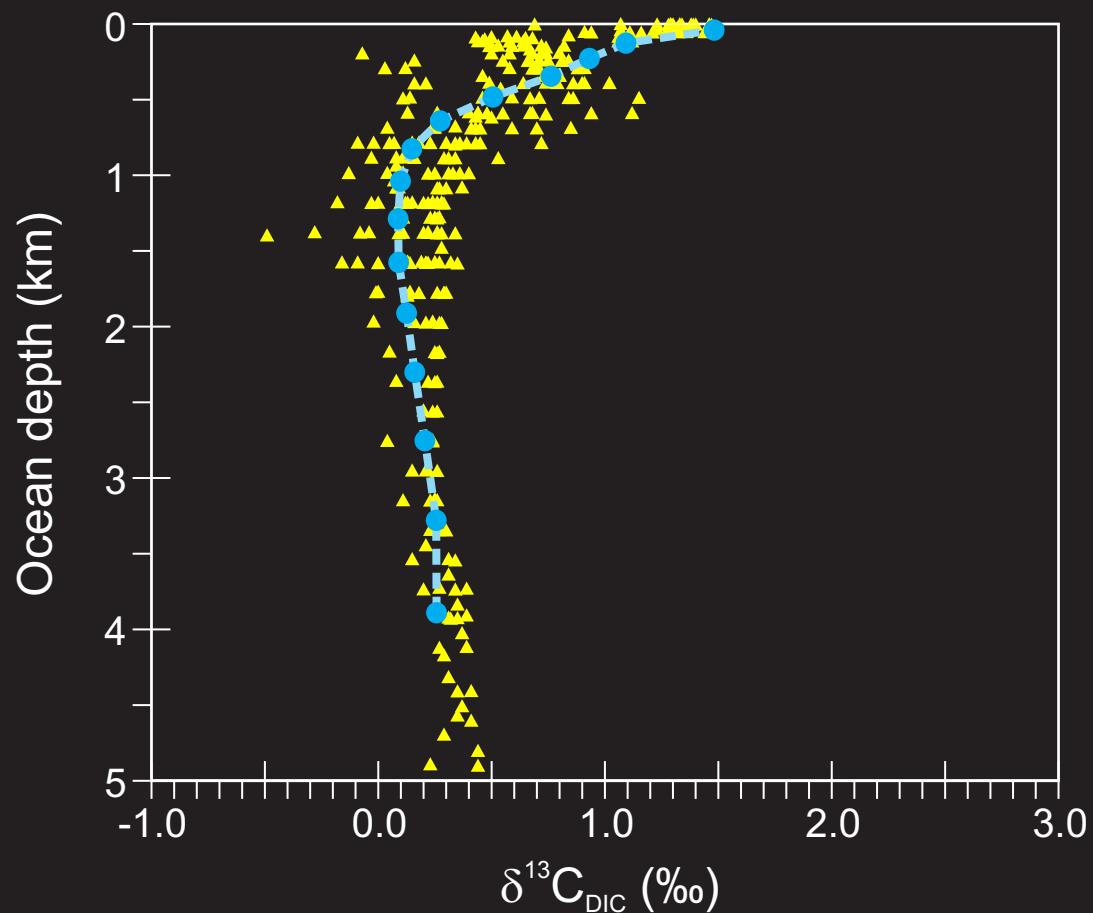
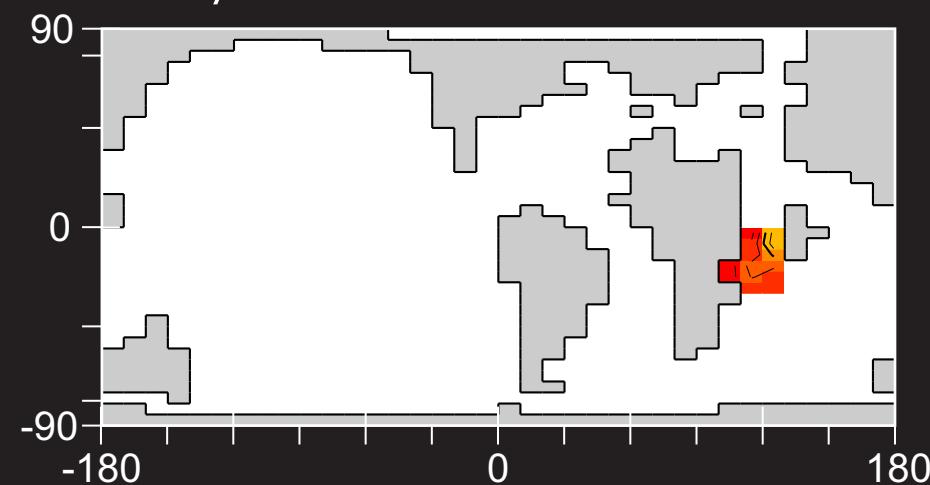
Planktic foraminiferal  $\delta^{13}\text{C}$  from early Eocene Tanzania



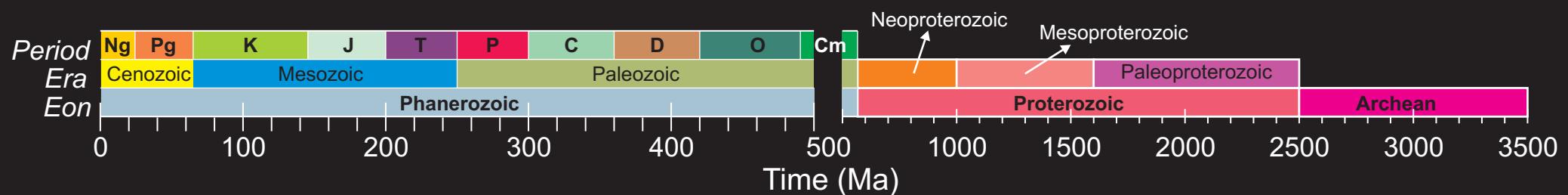
Open ocean  $\delta^{13}\text{C}_{\text{DIC}}$  adjacent to modern Tanzania



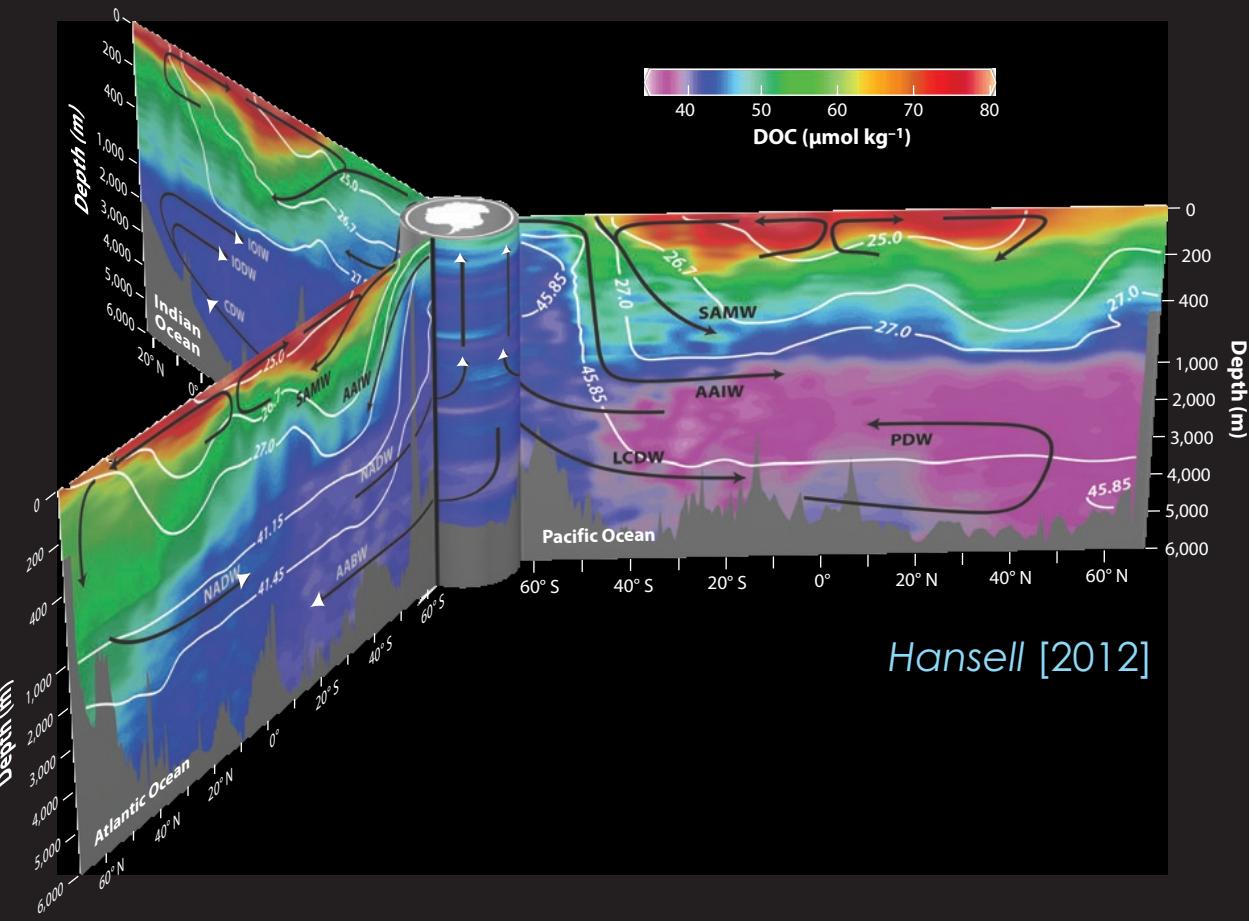
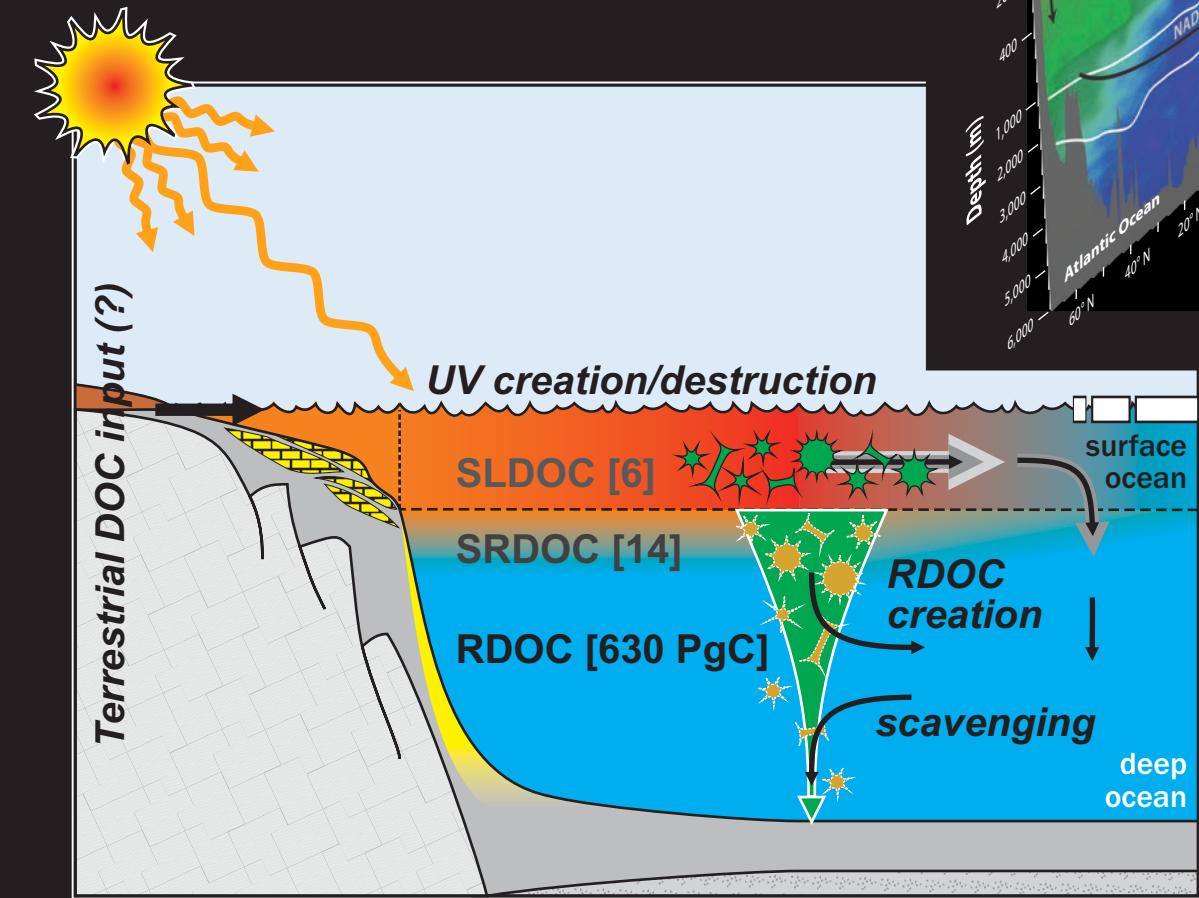
Planktic foraminiferal  $\delta^{13}\text{C}$  from early Eocene Tanzania



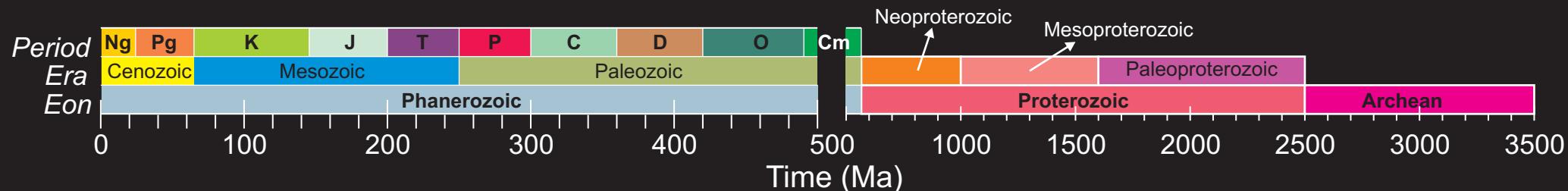
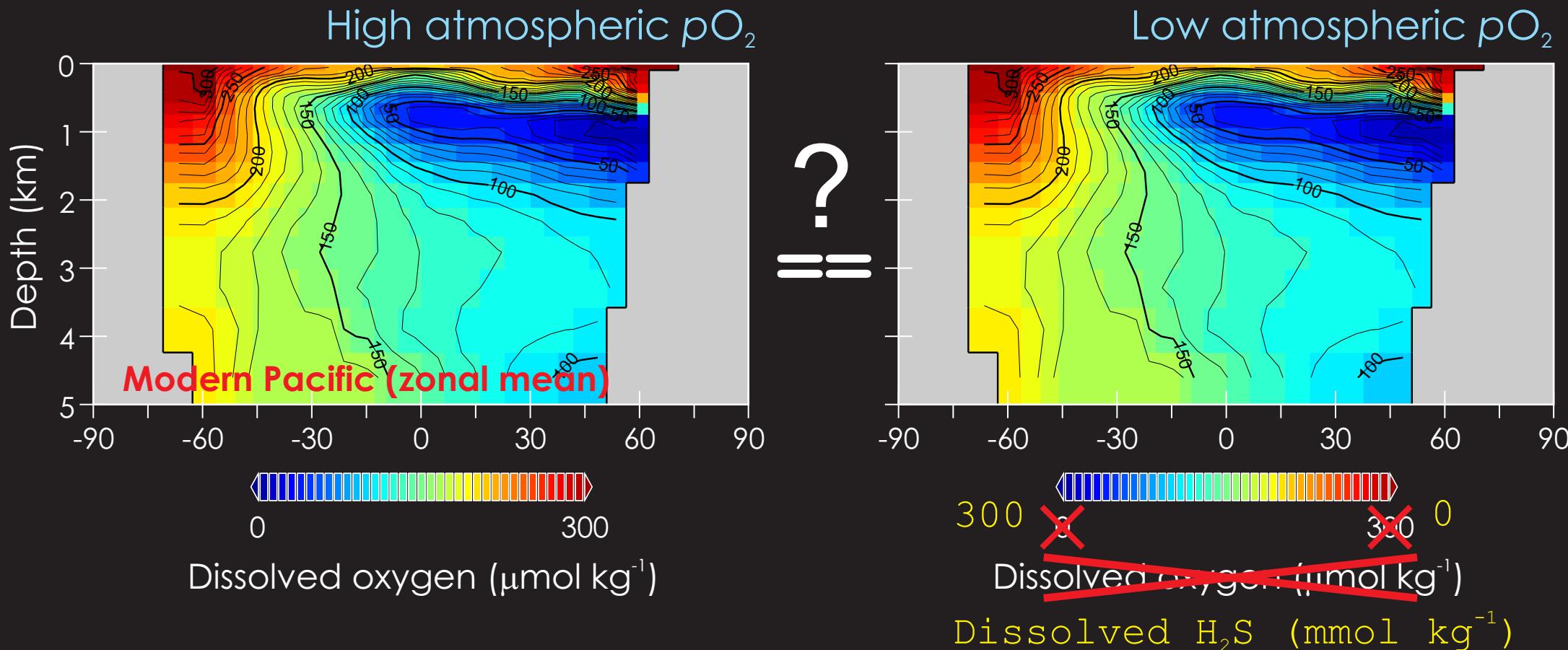
# *Evolution of the Biological Pump*



# Evolution of the Biological Pump: Dissolved organic matter



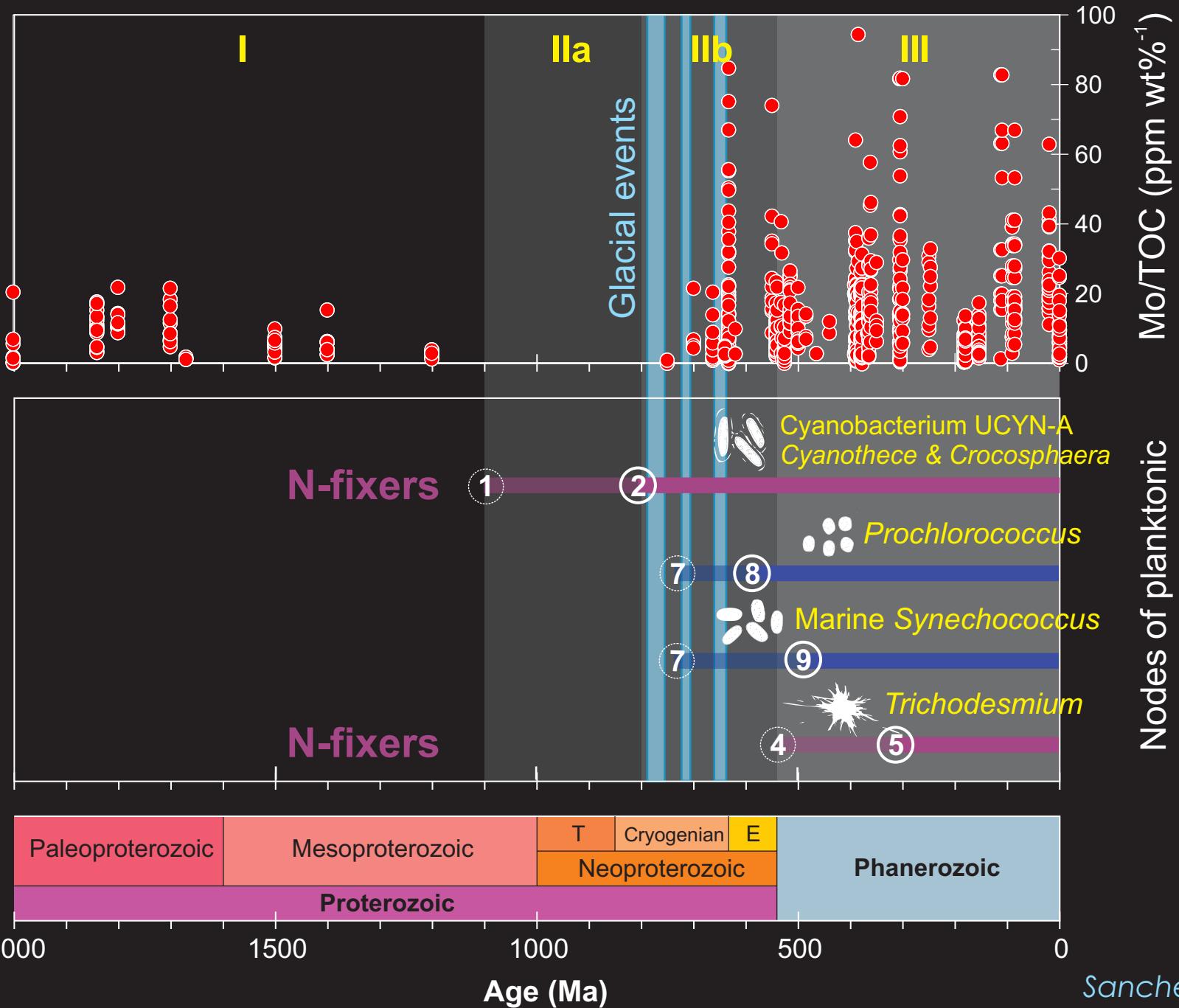
# *Evolution of the Biological Pump:* Dissolved organic matter



Low fixed N supply to the open ocean  
Low open ocean primary production

Transitional interval

High diversity of N fixers  
High primary production



*Thanks to:*

*Jamie Wilson & Steve Barker,  
Eleanor John, Paul Pearson [Cardiff]  
Patricia Sanchez-Baracaldo,  
Sandra Arndt, Daniela Schmidt [Bristol]*

*Ellen Thomas [Yale]*

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